1 Title: An Analytical Solution for the Run-out of Submarine Debris flows

2

3 Abstract

4 Submarine debris flows have a significant impact on offshore and coastal facilities. The 5 unique characteristics of submarine debris flows involve large mass movements and long 6 travel distances over very gentle slopes. To improve our insight and knowledge of the basic 7 mechanism behind submarine debris flows, an analytical model was derived for the mobility of 8 submarine debris flows. This model takes into account the mass change of debris flows 9 induced by deposition, stagnation pressure and the topography of the depositional area. One 10 case study on the Palos Verdes debris flow proves its ability to predict the run-out distance of 11 a submarine debris flow to a reasonable level of accuracy. On the gentle slopes, the 12 submarine debris flow progressively loses mass due to deposition, which in turn influence the 13 flow velocity. In addition, the results show that the slope angle and spreading angle of the 14 debris depositional zone play important roles in the sliding process.

15

16 Keywords

- 17 submarine debris flow, analytical solution, run-out distance, stagnation pressure
- 18

19 List of Notation

- 20 *P*: the stagnation pressure
- 21 ρ_w : the density of water
- 22 *v*: the flow velocity
- 23 τ_s : the friction force
- 24 F': the effective static pressure
- 25 ρ_f : the saturated bulk density of debris flow material
- 26 β_1 : the slope angle of escarpment
- 27 $\ddot{u}_{seismic}$: the seismic acceleration

1	θ:	the direction of seismic acceleration
2	b ₀ :	the initial width of debris flow
3	λ:	the spreading angle of debris deposition zone
4	β ₂ :	the slope angle of debris deposition zone
5	L:	the length of moving block
6	σ:	the density of particles
7	D:	the diameter of particles
8		

10 **1. Introduction**

11 Submarine debris flows are considered to be one of the most serious geohazards in offshore 12 and coastal areas. Due to the large affected area, even small-scale debris flows in coastal 13 areas can pose severe danger. Coastal infrastructure and populations, and offshore facilities 14 related to resource development and transport facilities, such as pipelines and communication 15 cables, are at risk from submarine debris flows, and must be designed to withstand their 16 impact. An example of submarine cable damage has been found in the Grand Banks debris 17 flow of 1929 where the debris flow and resulting turbidity current broke a series of submarine 18 cables nearly 600 km away from the beginning of the submarine debris flow (Hampton et al., 19 1996; Mason et al., 2006). Another case of infrastructure destruction occurred when 20 Hurricane Camille hit the Mississippi delta in 1969, causing a debris flow that damaged 21 several offshore drilling platforms (Locat and Lee, 2002).

22

Submarine debris flows are difficult to observe directly, and so it is difficult to obtain an insight into their behaviour. Many lab experiments have been carried out as small-scale analogues to investigate the roles of slurry properties and water content in debris flow dynamics and depositional structures (Mohrig et al., 1999; Marr et al., 2001; Breien et al., 2007; Boylan et al., 2010; Yin and Rui, 2017; Yin et al., 2018). However, these experiments are usually expensive, and the results are not easily transferable to large-scale conditions. To avoid these issues, a

1 considerable amount of numerical work has been done to use computer models to simulate 2 the failure or sliding of submarine debris flows (Imran et al., 2001; Marr et al., 2001; Gauer et 3 al., 2006; Zhu and Randolph, 2009; Steffen et al., 2008; Pudasaini, 2014; Soundararajan, 4 2015; Kafle et al. 2016). One early example is the Norem-Irgens-Schieldrop (NIS) model 5 proposed by Norem et al. (1990). The original purpose of the NIS model was to analyse snow 6 avalanches, however through analysis of the results it has been shown that the model is also 7 appropriate for use with submarine debris flows. Later Imran et al. (2001) proposed the BING 8 model, which is developed by incorporation of the Bingham, Herschel-Bulkley and Bilinear 9 rheologies. In the BING model the debris flow motion is considered as two coupled layers that 10 refer to a plug layer and a shear layer. Subsequent research (De Blasio et al., 2005) extended 11 the BING model to be able to calculate run-out behaviour with hydroplaning and shear-wetting. 12 Pudasaini (2012) presented a generalized two-phase debris flow model, which can 13 adequately simulate the complicated dynamics of submarine debris flows and related 14 phenomena. Based on this mode, Kattel et al. (2016) simulated a glacial lake outburst flood, 15 and showed that this model can be applied to simulate particle-fluid flow in conduits that is 16 important for decommissioning of the endangered reservoirs. In addition, Pudasaini (2014) 17 and Kafle et al. (2016) simulated two-phase debris impacting a reservoir, modeling the 18 landslide-induced tsunami, flow transformation, turbidity currents and sediment transports in 19 bathymetric slopes. On the other hand, Mergili et al. (2017, 2018) developed an open source, 20 efficient and high-resolution computational tool (r.avaflow) for routing mass flows from a 21 release area down to a deposition. This innovative tool showed many advantages over other 22 existing tools, such as (i) employment of a two-phase mixture model (Pudasaini, 2012); (ii) 23 capability of modelling complex process chains/interactions; (iii) modelling of multiple release 24 masses and/or hydrographs.

25

Although numerical models can offer a more accurate and detailed method of modelling flow dynamics, there are also some limitations to these numerical techniques, such as the amount of computing power needed and possible calculation errors. Other studies include empirical, semi-empirical and analytical modelling based on field observations and experimental data. Hampton et al. (1996) proposed that submarine debris flows can be described using rigid

1 body or continuum models. De Blasio et al. (2006) focused on the extraordinary mobility of 2 submarine debris flows, and concluded that the run-out ratios of vertical fall height to 3 horizontal travel distance may reach values as low as 0.05-0.01. Legros (2002) proposed 4 another closed solution for the mobility of long run-out debris flows, which indicates that the 5 debris flow spreading is essentially controlled by the volume, not the fall height. Boukpeti et al. 6 (2012) presented analytical models for different rheological models of the debris flow material, 7 but no solution was given and the run-out distance was not discussed. In addition, Pudasaini 8 and Miller (2013) presented a mass or volume dependent model for hypermobility of huge 9 landslides and avalanches. The model can be used to estimate the overrun area and volume 10 in terms of known mobility data.

11

12 Generally, analytical solutions for submarine debris flows require simplification of some 13 conditions by making assumptions on flow dynamics, since it is not possible to solve 14 complicated questions involving uncertain characteristics of soil/rock sediments, the 15 geometrical complexity of the sea bed, the interaction between a debris flow and the ambient 16 water, and the complicated mechanism behind sliding. However, these investigations are 17 widely considered to be important because they help to clarify the fundamental characteristics 18 of the phenomenon, and improve our insight and knowledge of the basic mechanism behind 19 submarine debris flows (Tinti and Bortolucci, 2000). This paper discusses and develops an 20 analytical approach to the run-out distance of submarine debris flows. This approach has 21 taken into account the impacts of mass change, stagnation pressure and depositional shape.

22

23 **2.** An analytical model for the submarine debris flow dynamics

Submarine debris flows describe the downward movement of a disturbed mass, after shear failure on the sliding surfaces. The disturbed mass travels down the slope in different manners according to the material attributes, topography and boundary conditions. Accordingly, the mobility behaviour of submarine debris flows can be divided into two categories. On a steep slope, the component of gravity in the downslope direction is usually larger than the friction between the moving mass and the base of the slope, and the mass of the whole submarine debris flow body remains constant without leaving deposits, as shown in

1 Figure 1. Similar assumptions could be found in the analytical model by Hürlimann et al. 2 (2000). On a gentle slope, due to the basal friction being higher than the component of gravity 3 force downslope, the material at the bottom of the displaced mass decelerates quickly until 4 completely stopped, but the upper part of the mass still moves downwards due to inertia. 5 Such depositional behaviours have also been simulated and shown with analytical models by 6 Legros (2002), Pudasaini and Kroener (2008), and Pudasaini (2011). This process indicates 7 deposition at the base of the debris flow, which thus progressively runs out of all of the 8 material. The mechanisms for submarine debris flows on steep and gentle slopes are quite 9 different, and hence two analytical models are needed for these two different cases.

10

11

- 12 Figure 1. Evolution of a submarine debris flow.
- 13

14 **2.1 Constant-mass model**

15 In the first model, the height and volume of debris flow are assumed to remain constant 16 through the flow process. Hence, this simplified model is called constant-mass model. Once 17 failure occurs, it is assumed that the debris flow comprises a homogenous mass which moves 18 over a fixed slope. On a steep slope, when the displaced mass has sufficient strength, the 19 submarine debris flow may avoid fragmentation and remain as one single body. Considering 20 the simplified process of a submarine debris flow as represented in Figure 2, several 21 assumptions are made: (a) the debris flow is considered to have a constant mass, and (b) the 22 height, length and width of the debris flow remain constant during the flow process (means it 23 does not deposit).

24

25

Figure 2. Schematic of a submarine debris flow (mass conservation) where Xmax = maximum
distance travelled.

28

The driving forces acting on the debris flow are proportional to the gravitational acceleration component parallel to the slope bed. Opposing these forces, as shown in Figure 2, are the

3 pressure is given by: $\mathbf{P} = \frac{1}{2}\rho_w v^2,$ 4 (1) 5 where ρ_w is the water density and v is the flow velocity of submarine debris. The friction 6 between the debris flow and the bed surface is assumed to follow classical Coulomb law, and 7 therefore can be presented as: $T = F' \cdot \tan \alpha$. 8 where T is the friction force between the debris flow and the bed surface, F' is the effective 9 10 11 12 effective static pressure, F' can be expressed as: $F' = (\rho_f - \rho_w)gV\cos\beta$, 13 14 15 16 gravity component of the force in the downslope direction with consideration of buoyancy is: $G'_{x} = (\rho_f - \rho_w) gV \sin \beta.$ 17 (4) 18 19 Accordingly, based on the conservation of energy, the debris flow mechanics can be 20 approximated by the following equation: 21 $dE = F_i \cdot dx$, (5) 22 where F_i is the net driving force given by (1), (2) and (4), x is the flow displacement, and E is 23 the kinetic energy. $\mathbf{F}_i = \mathbf{G}'_{\mathbf{x}} - \mathbf{T} - \mathbf{P} \cdot \mathbf{h}\mathbf{b},$ 24 $E = \frac{1}{2} \mathrm{m} v^2.$ 25 26 Hence, for constant m, equation (5) can be expanded as: $mv \frac{dv}{dx} = m'g \sin\beta - m'g \cos\beta \tan\alpha - \frac{1}{2}\rho_w v^2 hb,$ 27 (8) 28 where the mass of the sliding material, $m = \rho_f blh$, $m' = (\rho_f - \rho_w)blh$. This can be simplified to: $v \frac{\mathrm{d}v}{\mathrm{d}x} = g \frac{\rho_{\mathrm{f}} - \rho_{\mathrm{w}}}{\rho_{\mathrm{f}}} \sin \beta - \frac{\rho_{\mathrm{f}} - \rho_{\mathrm{w}}}{\rho_{\mathrm{f}}} g \cos \beta \tan \alpha - \frac{1}{2!} \frac{\rho_{\mathrm{w}}}{\rho_{\mathrm{f}}} v^{2}.$ 29 (9)

- (6)
- (7)

static pressure due the overburden with consideration of buoyancy, and α is the friction angle, considering the effects of water as a fluidized medium. (Pudasaini and Miller, 2013) The

friction with the slope bed and the drag force due to the surrounding fluid, caused by the

stagnation pressure (De Blasio, et.al, 2004). From Bernoulli's principle, the stagnation

1

2

(2)

(3)

where ρ_f is the density of the sliding material, h is its height, and β is the slope angle, V is the volume of the moving block, and V≈b·h·l (width, height and length of moving block). The

3

5

2 Equation (9) can be expressed in the form:

$$\frac{\mathrm{d}(\mathrm{v}^2)}{\mathrm{d}\mathrm{x}} = \mathrm{A} - \mathrm{B}\mathrm{v}^2,\tag{10}$$

4 where:

$$A = 2 \frac{\rho_f - \rho_w}{\rho_f} g \sin\beta - 2 \frac{\rho_f - \rho_w}{\rho_f} g \cos\beta \tan\alpha,$$
(11)

$$\mathsf{B} = \frac{\rho_{\mathsf{w}}}{\rho_{\mathsf{f}} \mathsf{l}}.$$
 (12)

7

6

8 This simplified constant-mass model is also mentioned by Hürlimann et al. (2000). The 9 difference is that Hürlimann's model is derived based on Newton's second law. In some cases 10 where the submarine debris flows are induced by the earthquake, the effect of the seismic 11 load on the moving mass needs to be included. This seismic load can be simplified as the 12 product of mass times constant seismic acceleration, $\ddot{u}_{seismic}$, in the direction of θ . This 13 seismic load is applied to the moving mass, as shown in Figure 2. Hence A can be expressed 14 as:

15
$$A = 2 \frac{\rho_f - \rho_w}{\rho_f} g \sin\beta + 2 \ddot{u}_{seismic} \cos(\theta - \beta) - 2 \frac{\rho_f - \rho_w}{\rho_f} g \cos\beta \tan\alpha - 2 \ddot{u}_{seismic} \sin(\theta - \beta) \tan\alpha$$

16 (13)

17

18 The flow displacement is obtained by integrating Equation (10) between X_0 and X_i which 19 correspond to the velocities v_0 and v_i , respectively:

20
$$-\frac{1}{B}\ln(A - Bv^2) = X_i - X_0 \qquad |_{v_0}^{v_i}$$
(14)

21

Accordingly, if the initial flow velocity is known, Equation (15) can be used to predict the valueof the velocity after a certain flow displacement.

24

25 2.2 Changing-mass model

During the run-out process the debris flow mass may change, due to deposition of material or accretion of material from the bed, which is the case in general during erosion and deposition process (Pudasaini and Fischer, 2016). Cannon and Savage (1988), Gassen and Cruden,

1 (1989), and Voight and Sousa (1994) showed that the impact of progressive changes to mass 2 due to deposition will increase the run-out distance of debris flow. Therefore, we may wonder 3 whether a debris flow that progressively loses mass due to deposition might not maintain a 4 higher velocity and travel further than a debris flow that moves and stops as a single block 5 (Legros, 2002; Pudasaini and Fischer 2016). According to this assumption, as presented in 6 Figure 3, mass loss occurs at some point where the debris flow starts to deposit material, and 7 only the upper part of debris proceed in the flow regime. Hence, a deposition model, which 8 could be used to simulate the erosion and spreading behaviour of submarine debris flow, 9 needs to be introduced. Pudasaini and Fischer (2016) indicated that the deposition models 10 can be divided into two categories: empirical and mechanical ones. Empirical model is always 11 calibrated by specific case and hence become case dependent. Process-based mechanical 12 models are based on the mass and momentum exchanges between the slope bed and sliding 13 debris, and the magnitude of erosion rate is shows a proportional relationship with the shear 14 stress difference between entraining and resisting stresses.(Iverson, 2012; Issler, 2014). 15 Within this framework, Legros (2002) proposed a deposition model, in which the thickness of 16 the upper part of debris which still proceed in the flow regime was expressed as,

18 where σ is the density of particles, D is the diameter, f_c is a positive function of the particle 19 concentration.

20

Assuming that $dv/dy = v/h_s$, where h_s is the typical thickness of the shearing zone, Legros (2002) simplified Equation (15) into

23
$$h = \frac{f_c \sigma D^2 \cos \alpha}{\rho_f g h_s^2} v^2 = B_1 v^2 , \qquad (16)$$

where B_1 is the material constant, which is dependent on granular concentration, grain diameter, density of particles and dispersive pressure. All the parameters in Equation 16 is assumed independent of v. On the other hand, equation (16) has also been derived analytically in Pudasaini (2011) in which the constant coefficient is explicitly expressed in terms of the mechanical and material parameters. It should be noted that mass loss is only represented in this model by a rebhnduction in flow height, and not in a real extent. (Legros,
 2002).

3

Figure 3. Schematic of a submarine debris flow run-out that includes transformation from the
constant-mass model to the non-constant-mass model.

6

9

7 By assuming that the area, (s = bl), of the moving debris flow does not vary with time (Legros, 8 2002), and that $m = blh\rho_f$, therefore:

$$dm = b l \rho_{\rm f} \cdot dh. \tag{17}$$

10 This assumption restricts the analytical model. However, Legros (2002) compared this 11 assumption and three cases of natural landslides, which shows the errors induced by this 12 assumption are acceptable. In addition, the results of some numerical simulations (Campbell 13 et al. 1995) also confirm this assumption.

$$\frac{d(\frac{1}{2}mv^2)}{dx} = G - T - Phb.$$
 (18)

17

16

18 Substituting equations (1), (2) and (4) into Equation (18):

19
$$\frac{v^2 dm}{2dx} + \frac{mv dv}{dx} = m' g \sin \beta - m' g \cdot \tan \alpha \cos \beta - \frac{1}{2} \rho_w v^2 hb.$$
(19)

20

The length of the moving mass, l, is also a variable based on the debris flow spreading in the horizontal direction. This debris flow spreading may be very complicated due to the complexity of the local topography, but it can be simplified as a radially spread deposit, as shown in Figure 4. The width of the moving mass can be written as $b(x) = b_0 + \lambda x$, where b_0 is the initial width and λ is the spreading angle. Hence, the length of the moving mass $l(x) = s/(b_0 + \lambda x)$.

28

29 (a) (b)

30 Figure 4. Plan view schematic of a debris flow on a gentle slope.

2 Equation (19) can be rewritten as:

3
$$\frac{v^2}{2h}\frac{dh}{dx} + v\frac{dv}{dx} = -\tan\alpha\cos\beta\frac{\rho_f - \rho_w}{\rho_f}g - \frac{1}{2}\frac{\rho_w}{\rho_f}v^2\frac{b_0 + \lambda x}{s} + \frac{\rho_f - \rho_w}{\rho_f}g\sin\beta.$$
(20)

It is essential to this model that the value of h changes with the debris flow, as described in
Equation 16. It is noted that B₁ is eliminated when substituting Equation (16) into Equation
(20), and hence it avoid the determination of those complicated parameters. Therefore
Equation (20) can be expressed in the form:

8
$$\frac{d(v^2)}{dx} + M + N(b_0 + \lambda x)v^2 = 0,$$
 (21)

9 where, $N = \frac{\rho_w}{2\rho_f s}$, $M = \tan \alpha \times \cos \beta \frac{\rho_f - \rho_w}{\rho_f} g - \frac{\rho_f - \rho_w}{\rho_f} g \sin \beta$, which can also include the seismic

10 load.

11

12 Therefore, a closed-form solution of Equation (21) can be expressed as:

13
$$v^{2} = \operatorname{Ce}^{-N\left(b_{0}x + \frac{1}{2}\lambda x^{2}\right)} + \operatorname{e}^{-N\left(b_{0}x + \frac{1}{2}\lambda x^{2}\right)} \frac{M\sqrt{\pi} \times \operatorname{i} \times \operatorname{e}^{-\frac{b_{0}^{2}N}{2}}\operatorname{derf}\left(\frac{\sqrt{2N}(b_{0} + x\lambda) \times \mathrm{i}}{2\sqrt{2}N\lambda}\right)}{\sqrt{2N\lambda}}, \qquad (22)$$

14 where C is a constant, which can be calculated by the initial condition and $i = \sqrt{-1}$.

15

16 The run-out distance X_{max} at which *v* becomes zero can be obtained implicitly by solving the 17 equation as:

18
$$\operatorname{Ce}^{-\operatorname{N}\left(b_{0}x+\frac{1}{2}\lambda x^{2}\right)}+\operatorname{e}^{-\operatorname{N}\left(b_{0}x+\frac{1}{2}\lambda x^{2}\right)}\frac{\operatorname{M}\sqrt{\pi}\times i\times \operatorname{e}^{-\frac{b_{0}^{2}\operatorname{N}}{2}}\operatorname{derf}\left(\frac{\sqrt{2\operatorname{N}\left(b_{0}+x\lambda\right)\times i}}{2\sqrt{2\operatorname{N}\lambda}}\right)}{\sqrt{2\operatorname{N}\lambda}}=0.$$
 (23)

On the other hand, if the mass-changing model (Equation 16) which describes the erosiondeposition behaviour is not considered, Equation (19) can be written as,

21
$$\frac{1}{2} \cdot \frac{d(v^2)}{dx} = -\tan\alpha\cos\beta\frac{\rho_f - \rho_w}{\rho_f}g + \frac{\rho_f - \rho_w}{\rho_f}g\sin\beta - \frac{\rho_w}{2\rho_f}\frac{b_0 + \lambda x}{s}v^2.$$
(24)

Accordingly, the final solution of debris flow velocity without considering mass-changing can be rewritten. So, essentially the mapping $N \rightarrow N/2$ was from none-mass-changing solution to mass-changing solution,

25
$$v^{2} = Ce^{-2N\left(b_{0}x + \frac{1}{2}\lambda x^{2}\right)} + e^{-2N\left(b_{0}x + \frac{1}{2}\lambda x^{2}\right)} \frac{M\sqrt{\pi} \times i \times e^{-b_{0}^{2}\frac{N}{\lambda}} \operatorname{erf}\left(\frac{\sqrt{N}(b_{0} + x\lambda) \times i}{\sqrt{\lambda}}\right)}{\sqrt{N\lambda}}.$$
 (25)

3. Application of analytical models

2 The proposed analytical model was used to analyse a well-known event: the Palos Verdes 3 debris flow (Hampton et al., 1996; Locat and Lee, 2002; Locat et al., 2004), located on the 4 continental slope near Los Angeles. Based on seismic reflection logs of the local morphology, 5 the feature of the sea floor lying at the base of the escarpment was recognised as a 6 submarine rock avalanche (Gorsline et al. 1984). The debris avalanche deposit at the toe of 7 the escarpment resulted from a failure that took place along the upper part of the escarpment. 8 The morphology of the escarpment is described in Figure 5. The slope of the escarpment 9 varies between 10° and 17°. At the toe of the escarpment, the debris avalanche deposit 10 spread for a distance of about 8 km, over a slope varying between 1.5° and 2°. The Palos 11 Verdes debris flow process can be divided into two different stages, according to the material 12 attributes and topography. The first stage refers to a constant-mass debris flow process on 13 the escarpment (failure plane). The debris flow is simplified as a whole block with constant 14 mass for the whole sliding process. The sliding distance is about 1km, as shown in Figure 5. 15 Due to the variety of slope, the calculation is divided into 10 steps, referring to 10 sections of 16 escarpment in Figure 5(b). In this process, the spreading angle λ is assumed as 0°, and 17 hence the width b of moving block always equals to the initial value 1000m. Other parameters 18 are listed in Table 1 and Table 2. The initial velocity v_0 is 0m/s, and the final velocity in 19 section one v_{0.1km} (velocity at distance of 0.1km) can be calculated by Equation 14. This 20 velocity v_{0.1km} would be the initial velocity in the next step. Repeating this calculation could 21 get the final velocity of constant-mass debris flow $v_{1\rm km}$, defined as the velocity at the toe of 22 the escarpment and initial velocity of deposit zone. The second stage refers to the debris flow 23 in which debris flow was deposited in the debris deposit zone (section 11 in Figure 5(b)). The 24 debris flow is simplified as a whole block as well, but with varying mass. The final sliding 25 distance x_{final} can be calculated by solving Equation (23). The change in the slope of deposit 26 zone is very limited, hence the deposit zone is treated as one whole section. In this model, 27 the the basal topography changes with the changing mass is not considered, which may 28 influence the entire flow dynamics. Hence, the future work will focus on the implementation of 29 the full mechanical models (Pudasaini and Fischer, 2016; Mergili et al., 2017, 2018).

(a)

(b)

2	Figure 5. San Pedro Escarpment and the Palos Verdes avalanche deposit: (a) plan view
3	(Google map, 2017); (b) side view (after Locat et al., 2004).

4

5

6

Table 1 Geometry of escarpment for the Palos Verdes debris flow.

7 Locat et al. (2004) conducted a numerical analysis of the mobility of the Palos Verdes debris 8 avalanche. The analysis of the failure stage indicated that the debris avalanche was caused 9 by a major earthquake with a magnitude around 7 on the Richter scale, corresponding to a 10 seismic acceleration of 0.3-0.4 g in the downslope direction, where g is the gravitational 11 acceleration. Some of the other parameters from Locat et al. (2004) for the Palos Verdes 12 debris flow were also used in the analytical model, and are listed in Table 2. The variety in the 13 value of average slope angle of escarpment, slope and spreading angle of debris deposition 14 zone is used for parametric study.

- 15
- 16

Table 2 Parameters for analysis of the Palos Verdes debris flow.

17

18 The above values listed in Table 2 were applied to the analysis. The results are given in 19 Figure 6, with the dash line indicating the most probable values for the real situation about 20 8km. (Locat et al., 2004) The figure shows that most of the results are close to the real value. 21 Most of the relative errors are below 20%. Hence it demonstrates the ability of the analytical 22 solution to predict the run-out distance of a submarine debris flow to a reasonable level of 23 accuracy. Figure 6(a) shows the prediction of the run-out distance with varying slope angles 24 for the escarpment and debris deposition zone. It is shown that the run-out distance increases 25 with increase in either the average slope angle of the escarpment, β_1 , or the angle of the 26 debris deposition zone, β_2 . When $\beta_2 = 2^\circ$, variations of β_1 from 10° to 17° lead to an increase 27 in the run-out distance from about 7.9 km to 9.5 km. On the other hand, as β_2 increases from 28 1.5° to 2°, the average change in the run-out distance of all the cases is about 2 km. 29 Therefore, this result indicates that β_2 has a greater influence on the overall run-out distance

1 than β_1 . Figure 6(b) shows the prediction of the run-out distance with varying escarpment 2 slope angles and spreading angles of the debris deposition zone. The effect of the spreading 3 angle of debris deposition zone on the run-out distance was investigated, showing an inverse 4 relationship between the run-out distance and the spreading angle of the debris deposition 5 zone. A larger spreading angle, λ , causes a larger width of the moving mass, and hence a 6 comparatively larger stagnation pressure applied to the front of the moving mass. Therefore, 7 the overall run-out distance decreases when the spreading angle, λ , increases. In addition, in 8 order to estimate the effect of mass-changing model, which is described in Equation (16), a 9 sensitivity analysis has been performed as shown in Figure 6(c). Equation (25) describes this 10 simplified model. It is shown that the calculated run-out distance decreased about 40% in all 11 cases when the mass-changing model is not considered. Therefore, it can be concluded that 12 mass-changing and deposition play important roles in the extraordinary mobility of submarine 13 debris flow. However, the present model cannot explain the mechanism for why reduction of 14 sliding mass results in increased mobility. The two-phase mechanical erosion-deposition 15 model by Pudasaini and Fischer (2016) proved that erosion enhances flow mobility, while 16 deposition reduces mobility.

- 17
- 18

19 20 (a)

Figure 6. Calculated run-out distance: (a) with varying slope angles of the escarpment and the debris deposition zone; (b) with varying escarpment slope angles and spreading angles of the debris deposition zone; (c) with and without considering the mass changing model.

(C)

(b)

24

25

26 4. Conclusions

This paper has presented analytical solutions for determining the run-out distance of a submarine debris flow, with the aim of exploring the sensitivity of the results to different geomechanical attributes and environmental factors. The following conclusions are derived:

(1) The analytical solutions proposed in this paper consider the erosion and spreading
 behaviour of submarine debris flows, and hence are able to predict the run-out distance of
 submarine debris flows to a reasonable level of accuracy.

4 (2) If the debris flow material has sufficient shear strength, the moving block may avoid 5 fragmentation and remain as one single body on a steep sliding surface. However, debris 6 deposition occurs on gentle slopes. The debris flow material progressively loses mass due to 7 deposition, which has a large influence on the flow velocity. It is shown that the calculated 8 run-out distance decreased about 40% when neglecting the mass-changing model. But its 9 mechanical significance still needs future research to be prove, because it contradicts with the 10 mechanically derived model by Pudasaini and Fischer (2016) which clearly shows that only 11 relatively increased friction leads to deposition.

(3) Water is an important factor for submarine debris flows. It not only affects the friction angle
as a fluidising medium, but also applies stagnation pressure to the moving block. The
additional resistance from water leads to a decrease in velocity, and hence a decrease of the
final run-out distance of the submarine debris flow.

(4) The run-out distance of submarine is primarily controlled by the local slope of the
depositional zone. This affect is much larger than that from the angle of the escarpment slope.
(5) There is an inverse relationship between the run-out distance and the spreading angle of
the debris deposition zone, since the increase in the stagnation pressure induced by the
spreading of the deposition zone.

21

22 Acknowledgement

²³ This research work is part of the activities of the Schofield Centre at University of Cambridge.

²⁴ The work presented in the paper is part of the Modelling of Mudslide Runout Project which

²⁵ was a collaborative project between University of Cambridge and BP.

26

We thank Dr. Stuart Haigh, (University of Cambridge) and Professor Kenichi Soga,
 (UC Berkeley) for comments that greatly improved the research results, and Mr Takaaki
 Kobayashi for the assistance with the whole project. We would also like to show our gratitude

to Dr. Paul Dimmock (BP) for fully supporting this project and sharing his pearls of wisdom
 with us during the collaboration.

3

4 Reference

- Boukpeti, N., White, D. J., Randolph, M. F., & Low, H. E. (2012). Strength of fijkne-grained
 soils at the solid–fluid transition. Geotechnique 62, No. 3, pp. 213–226.
- Boylan, N., Gaudin, C., White, D.J., & Randolph, M. F. (2010). Modelling of submarine slides
 in the geotechnical centrifuge. 7th Int. Conf. on Physical Modelling in Geotechnics
 (ICPMG). Zurich. 1095-1100.
- Breien, H., Pagliardi, M., Blasio, F., Issler, D., & Elverhui, A. (2007). Experimental studies of
 subaqueous vs. subaerial debris flows-velocity characteristics as a function of the
 ambient fluid. Submarine Mass Movements and Their Consequences, pp. 101--110.
- Campbell, C. S., Cleary, P. W., & Hopkins, M. (1995). Large-scale landslide simulations:
 Global deformation, velocities and basal friction. Journal of Geophysical Research:
 Solid Earth, 100(B5), 8267-8283.
- Cannon, S. H., & Savage, W. Z. (1988). A mass-change model for the estimation of debrisflow runout. The Journal of Geology, 221-227.
- De Blasio, F. V., Engvik, L., Harbitz, C. B., & Elverhøi, A. (2004). Hydroplaning and
 submarine debris flows. Journal of Geophysical Research: Oceans (1978–
 20 2012), 109(C1).
- De Blasio, F. V., Elverhøi, A., Issler, D., Harbitz, C. B., Bryn, P., & Lien, R. (2005). On the
 dynamics of subaqueous clay rich gravity mass flows—the giant Storegga slide,
 Norway. Marine and Petroleum Geology, 22(1), 179-186.
- De Blasio, F. V., Elverhoi, A., Engvik, L. E., Issler, D., Gauer, P., & Harbitz, C. (2006).
 Understanding the high mobility of subaqueous debris flows.NORSK GEOLOGISK
 TIDSSKRIFT, 86(3), 275.
- Gassen, W. V., & Cruden, D. M. (1989). Momentum transfer and friction in the debris of rock
 avalanches. Canadian Geotechnical Journal, 26(4), 623-628.

1	Gauer, P., Elverhoi, A., Issler, D., & De Blasio, F. V. (2006). On numerical simulations of
2	subaqueous slides: back-calculations of laboratory experiments of clay-rich slides.
3	NORSK GEOLOGISK TIDSSKRIFT, 86(3), 295.
4	GOOGLE MAPS, 2017. Map of San Pedro Escarpment: Undersea Features. [online]. Google.
5	Available from: <u>http://www.geographic.org/geographic_names/name.php?uni=-</u>
6	241460&fid=6437&c=undersea_features [Accessed 1May 2017].
7	Gorsline, D. S., Kolpack, R. L., Karl, H. A., Drake, D. E., Thornton, S. E., Schwalbach, J. R.,
8	& Fleischer, P. (1984). Studies of fine-grained sediment transport processes and
9	products in the California Continental Borderland. Geological Society, London,
10	Special Publications, 15(1), 395-415.
11	Hampton, M., Lee, H., & Locat, J. (1996). Submarine landslides. Reviews of Geophysics
12	34(1), pp. 3359.
13	Hürlimann, M., Garcia-Piera, J. O., & Ledesma, A. (2000). Causes and mobility of large
14	volcanic landslides: application to Tenerife, Canary Islands. Journal of Volcanology
15	and Geothermal Research, 103(1-4), 121-134.
16	Imran, J., Harff, P. and Parker, G. (2001). A numerical model of submarine debris flow with
17	graphical user interface. Computers & Geosciences, 27: 717-729.
18	Issler, D. (2014). Dynamically consistent entrainment laws for depth-averaged avalanche
19	models. Journal of Fluid Mechanics, 759, 701-738. Iverson, R. M. (1997). The physics
20	of debris flows. Reviews of geophysics, 35(3), 245-296.
21	Iverson, R. M. (2012). Elementary theory of bed - sediment entrainment by debris flows and
22	avalanches. Journal of Geophysical Research: Earth Surface, 117(F3).
23	Kafle, J., Pokhrel, P. R., Khattri, K. B., Kattel, P., Tuladhar, B. M., & Pudasaini, S. P. (2016).
24	Landslide-generated tsunami and particle transport in mountain lakes and
25	reservoirs. Annals of Glaciology, 57(71), 232-244.
26	Kattel, P., Khattri, K. B., Pokhrel, P. R., Kafle, J., Tuladhar, B. M., & Pudasaini, S. P. (2016).
27	Simulating glacial lake outburst floods with a two-phase mass flow model. Annals of
28	Glaciology, 57(71), 349-358.
29	Legros, F. (2002). The mobility of long-runout landslides. Engineering Geology, 63(3), 301-
30	331.

- Locat, J., & Lee, H. J. (2002). Submarine landslides: advances and challenges. Canadian
 Geotechnical Journal, 39(1), 193-212.
- Locat, J., Lee, H. J., Locat, P., & Imran, J. (2004). Numerical analysis of the mobility of the
 Palos Verdes debris avalanche, California, and its implication for the generation of
 tsunamis. Marine Geology, 203(3), 269-280.
- Marr, J., Harff, P., Shanmugam, G., & Parker, G. (2001), Experiments on subaqueous sandy
 gravity flows: The role of clay and water content in flow dynamics and depositional
 structures. Geological Society of America Bulletin 113(11), pp. 1377.
- 9 Masson, D., Harbitz, C., Wynn, R., Pederson, G., & Lovholt, F. (2006). Submarine
 10 landslides: processes, triggers and hazard protection. Philosophical Transactions of
 11 the Royal Society, 364, 2009-2039.
- Mergili, M., Jan-Thomas, F., Krenn, J., & Pudasaini, S. P. (2017). r. avaflow v1, an advanced
 open-source computational framework for the propagation and interaction of two phase mass flows. Geoscientific Model Development, 10(2), 553.
- 15 Mergili, M., Emmer, A., Juřicová, A., Cochachin, A., Fischer, J. T., Huggel, C., & Pudasaini, S.
- P. (2018). How well can we simulate complex hydro-geomorphic process chains?
 The 2012 multi-lake outburst flood in the Santa Cruz Valley (Cordillera Blanca,
 Perú). Earth Surface Processes and Landforms.
- Mohrig, D., Elverhøi, A., & Parker, G. (1999). Experiments on the relative mobility of muddy
 subaqueous and subaerial debris flows, and their capacity to remobilize antecedent
 deposits. Marine Geology, 154(1), 117-129.
- Marr, J. G., Elverhøi, A., Harbitz, C., Imran, J., & Harff, P. (2002). Numerical simulation of
 mud-rich subaqueous debris flows on the glacially active margins of the Svalbard–
 Barents Sea. Marine Geology, 188(3), 351-364.
- Norem, H., Locat, J., & Schieldrop, B. (1990). An approach to the physics and the modeling of
 submarine flowslides. *Marine Georesources & Geotechnology* 9(2), 93—111.
- 27 Pudasaini, S. P., & Kröner, C. (2008). Shock waves in rapid flows of dense granular materials:
- 28 Theoretical predictions and experimental results. *Physical Review E*, 78(4), 041308.
- Pudasaini, S. P. (2011). Some exact solutions for debris and avalanche flows. *Physics of Fluids*, *23*(4), 043301.

1	Pudasaini, S. P. (2012). A general two - phase debris flow model. Journal of Geophysical
2	Research: Earth Surface, 117(F3).
3	Pudasaini, S. P., & Miller, S. A. (2013). The hypermobility of huge landslides and
4	avalanches. Engineering Geology, 157, 124-132.
5	Pudasaini, S. P. (2014). Dynamics of submarine debris flow and tsunami. Acta
6	Mechanica, 225(8), 2423-2434.
7	Pudasaini, S. P., & Fischer, J. T. (2016). A mechanical erosion model for two-phase mass
8	flows. arXiv preprint arXiv:1610.01806.
9	Soundararajan, K. K. (2015). Multi-scale multiphase modelling of granular flows. PhD Thesis,
10	University of Cambridge, UK.
11	Steffen, M., Kirby, R. M., & Berzins, M. (2008). Analysis and reduction of quadrature errors in
12	the material point method (MPM). International Journal for Numerical Methods in
13	Engineering, 76(6), 922-948.
14	Tinti, S., & Bortolucci, E. (2000). Analytical investigation on tsunamis generated by submarine
15	slides.
16	Voight, B., & Sousa, J. (1994). Lessons from Ontake-san: a comparative analysis of debris
17	avalanche dynamics. Engineering Geology, 38(3), 261-297.
18	Yin, M., & Rui, Y. (2017). Laboratory study on submarine debris flow. Marine Georesources &
19	Geotechnology, 1-9.
20	Yin, M., Rui, Y., & Xue, Y. (2017). Centrifuge study on the runout distance of submarine
21	debris flows. Marine Georesources & Geotechnology, 1-11.
22	Zhu, H., & Randolph, M. F. (2009). Large deformation finite-element analysis of submarine
23	landslide interaction with embedded pipelines. International Journal of Geomechanics.
24	
25	Figure captions
26	Figure 1 Evolution of a submarine debris flow
27	Figure 2 Schematic of a slide (mass conservation) where Xmax = maximum distance
28	travelled
29	Figure 3 Schematic of a submarine debris flow run-out that includes transformation from the
30	constant-mass model to the non-constant-mass model

1	Figure 4 Disputient este stide en e sertie siere
1	Figure 4 Plan view schematic of a slide on a gentie slope
2	Figure 5. San Pedro Escarpment and the Palos Verdes avalanche deposit: (a) plan view
3	(Google map, 2017); (b) side view (after Locat et al., 2004)
4	Figure 6 Calculated run-out distance: (a) with varying slope angles of the escarpment and the
5	debris deposition zone; (b) with varying escarpment slope angles and spreading angles of the
6	debris deposition zone; (c) with and without considering the mass changing model
7	
8	Table captions
9	Table 1 Geometry of escarpment
10	Table 2 Parameters for analysis of the Palos Verdes slide
11	
12	
13	
14	
15	
16	
17	
18	
19	
20	
21	
22	
23	
24	
25	





Figure 3. Schematic of a submarine debris flow run-out that includes transformation from the constant-mass model to the non-constant-mass model











Figure 6. Calculated run-out distance: (a) with varying slope angles of the
escarpment and the debris deposition zone; (b) with varying escarpment slope
angles and spreading angles of the debris deposition zone; (c) with and without
considering the mass changing model



Table 1 Geometry of escarpment for the Palos Verdes debris flow

Section ID	Length	Slope Angle	Section ID	Length	Slope Angle
1	0.1km	13.5°	7	0.1km	15.4°
2	0.1km	15.4°	8	0.1km	17°
3	0.1km	15.9°	9	0.1km	11.3°
4	0.1km	16.3°	10	0.1km	10°
5	0.1km	16.7°	11	unknown	1.75°
6	0.1km	16.2°			

Table 2 Parameters for analysis of the Palos Verdes debris flow

Unit weight of debris-flow material (ρ_{f})	25 kN/m ³	Acceleration ($\dot{u}_{seismic}$)	0.35 g
Friction angle (α)	30°	Acceleration direction (θ)	10°
Average slope angle of escarpment (β_1)	10–17°	Length of escarpment	1000 m
Slope angle of debris deposition zone (β_2)	1.5–2°	Spreading angle of debris deposition zone (λ)	8–16°
Initial length of moving block (I)	1000 m	Initial width of moving block $(\mathbf{b_0})$	1000 m