

1 **Hydrologic Risk Analysis in the Yangtze River basin through Coupling Gaussian**
2 **Mixtures into Copulas**

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25 **Abstract:**

26 In this study, a bivariate hydrologic risk framework is proposed through coupling
27 Gaussian mixtures into copulas, leading to a coupled GMM-copula method. In the
28 coupled GMM-Copula method, the marginal distributions of flood peak, volume and
29 duration are quantified through Gaussian mixture models and the joint probability
30 distributions of flood peak-volume, peak-duration and volume duration are
31 established through copulas. The bivariate hydrologic risk is then derived based on
32 the joint return period of flood variable pairs. The proposed method is applied to the
33 risk analysis for the Yichang station on the main stream of the Yangtze River, China.
34 The results indicate that (i) the bivariate risk for flood peak-volume would keep
35 constant for the flood volume less than $1.0 \times 10^5 \text{ m}^3/\text{s day}$, but present a significant
36 decreasing trend for the flood volume larger than $1.7 \times 10^5 \text{ m}^3/\text{s day}$; (ii) the bivariate
37 risk for flood peak-duration would not change significantly for the flood duration less
38 than 8 days, and then decrease significantly as duration value become larger. The
39 probability density functions (pdfs) of the flood volume and duration conditional on
40 flood peak can also be generated through the fitted copulas. The results indicate that
41 the conditional pdfs of flood volume and duration follow bimodal distributions, with
42 the occurrence frequency of the first vertex decreasing and the latter one increasing as
43 the increase of flood peak. The obtained conclusions from the bivariate hydrologic
44 analysis can provide decision support for flood control and mitigation.

45

46 **Keywords:** Flood risk; Copula; Flood frequency analysis; Distribution; Conditional
47 distribution; Gaussian mixture model.

48

49 **1. Introduction**

50 Extreme hydrologic events, such as floods, droughts and storms, have been
51 leading to extensive property losses in recent decades. Specifically, floods have
52 become one of the most common natural disasters, posing significant risks to human
53 beings and environment [22, 31, 41, 67, 71-72, 81]. Hydrological frequency analysis
54 procedures are widely adopted to estimate the occurrence probabilities of floods,
55 providing decision support for many water resources management practices, such as
56 reservoir management, dam design and flood insurance studies [6, 16-20, 43, 45, 82-
57 85]. Moreover, a flood is associated with multidimensional characteristics.
58 Consequently, flood frequency analysis under consideration of multiple flood
59 variables would be desired to provide a full screen for a flood.

60 Copula functions, in recent years, have been widely used for multivariate
61 hydrologic modeling, such as multivariate flood frequency analysis [7, 18, 22, 26-27,
62 65, 78-79], drought assessments [14, 35, 42, 59, 60, 64], storm or rainfall dependence
63 analysis [2-4, 66], streamflow simulation [32, 39-40, 58]. De Michele and Salvador
64 [13] initially introduced the concept of copulas into hydrological simulation, which
65 described the dependence between storm duration and average rainfall intensity by
66 means of a suitable 2-Copula. Salvadori and De Michele [51] characterized the
67 dependence between storm duration and intensity via a suitable 2-Copula with the
68 marginal distributions endowed with Generalized Pareto laws. Recently, Salvador and
69 De Michele [52] conducted multivariate real-time assessment of droughts via copula-
70 based multi-site Hazard Trajectories and Fans. The main advantage of copula
71 functions over classical bivariate frequency analyses is that the selection of marginal
72 distributions and multivariate dependence modelling are two separate processes,

73 giving additional flexibility to the practitioner in choosing different marginal and joint
74 probability functions [24, 36, 65, 78]. Consequently, the selection of marginal
75 distributions would definitely impact the performance of the copula in modelling
76 multivariate hydrologic simulation.

77 In multivariate hydrologic frequency analysis through copula functions, the flood
78 variables under consideration include: the annual maximum peak discharges, and the
79 associated hydrograph volumes and durations. Consequently, the distributions for
80 modelling these flood variables would be various. For example, for modelling the
81 annual maximum flood series, the used distributions over the world include extreme
82 value type 1 (EV1), general extreme value (GEV), extreme value type 2 (EV2), two
83 component extreme value, normal, lognormal (LN), Pearson type 3 (P3), Log Pearson
84 type 3 (LP3), Gamma, exponential, Weibull, generalised Pareto and Wakeby
85 distributions [5, 12]. Previous studies have shown that, in modelling multivariate
86 flood frequency through copula functions, the marginal distributions of peak, volume
87 and duration were different at different sites. For example, Sraj et al. [65] took
88 bivariate flood frequency analysis using copula function for the Litija station on the
89 Sava River, in which log-Pearson 3 distribution was chosen for modelling discharge
90 peaks and hydrograph durations, and the Pearson 3 distribution was selected for
91 hydrograph volumes. Reddy and Ganguli [48] applied Archimedean copulas for
92 bivariate flood frequency analysis, where the normal kernel density function was used
93 for quantifying the distributions of peak flow and duration, and quadratic kernel
94 density function was applied for volume.

95 A Gaussian mixture model (GMM) is a mixture statistical model of a finite
96 number of Gaussian distributions with unknown parameters. It is a semiparametric
97 probability density function expressed as a weighted sum of Gaussian component

98 densities, and all samples are assumed to be generated from this mixture model.
99 GMMs are commonly used to model the probability distributions of continuous
100 measurements or features in a biometric system, such as vocal-tract related spectral
101 features in a speaker recognition system [50]. The finite Gaussian mixture model can
102 theoretically approximate any continuous distribution very closely if properly given a
103 sufficient number of components [34]. Several research works have been reported to
104 apply mixed distribution models to analyze hydrological and environmental data [15,
105 21, 34, 61-62, 68]. For example, Yue et al. [76] proposed a Gumbel mixed model for
106 flood frequency analysis. Singh et al. [62] proposed a mixed distribution method for
107 nonidentically distributed hydrologic flood data. He [34] applied the GMM for
108 analyzing the multiply censored environmental data. Although GMM and other mixed
109 model methods have been widely proposed to model the water and environmental
110 samples, these proposed methods have some limitations in practical multivariate flood
111 risk analysis. For instance, the Gumbel mixed distribution proposed by Yue et al. [76]
112 can only applied to positively correlated random variables with the correlation
113 coefficient less or equal to $2/3$ [77].

114 As an extension of previous research, this paper aims to couple the GMM into
115 copulas, leading to a coupled GMM-copula method for multivariate hydrologic risk
116 analysis. The advantages of the proposed method are that i) the GMM can provide
117 good estimations for the marginal distributions and ii) the copula method can relax the
118 assumptions in previous mixed models such as same type distribution, correlation
119 restriction [78]. Moreover, an integrated multivariate risk indicator is proposed to
120 reveal significance of effects from persisting high risk levels due to impacts from
121 multiple interactive flood variables. Such an analysis will be based on provision of the
122 coupled GMM-copula method. Finally, the conditional probability density

123 distributions (pdfs) of flood volume and duration under peak flows with different
124 return periods will be characterized, intending to explore potential control and
125 management practices once a flood has occurred. The proposed method will be
126 applied to the Yangtze River (Chang Jiang), China

127

128 **2. Methodology**

129 **2.1 Gaussian Mixture Model**

130 The mixture model is a useful tool for density estimation, and can be viewed as a
131 kind of kernel method [33]. Mixture models can use any component densities but the
132 Gaussian mixture model (GMM) is the most popular [33]. The probability density
133 function of a Gaussian mixture model is expressed by a weighted sum of M-
134 component Gaussian probability densities as given below:

$$135 \quad p(x) = \sum_{j=1}^M \alpha_j N_j(x; \mu_j, \sigma_j) \quad (1)$$

136 where x are one-dimensional measurement samples; α_j ($j = 1, 2, \dots, M$) denote the
137 mixture weights; $N_j(x; \mu_j, \sigma_j)$ ($i = 1, 2, \dots, M$) are the component Gaussian densities,
138 which can be expressed as:

$$139 \quad N_j(x; \mu_j, \sigma_j) = \frac{1}{\sqrt{2\pi}\sigma_j} \exp\left(-\frac{(x - \mu_j)^2}{2\sigma_j^2}\right) \quad (2)$$

140 where μ_j and σ_j respectively denote the mean and standard deviation for the j^{th}

141 Gaussian distribution model. The weights α_j are nonnegative and must satisfy

142 $\sum_{j=1}^M \alpha_j = 1$. The GMM has two main advantages in practical applications in many

143 engineering fields: (i) it can sufficiently approximate a broad class of distribution

144 functions encountered in practice, if an appropriate size of components are given in

145 the mixture; (ii) the form of the GMM simplifies the derivation of the subsequent

146 estimation method and avoids the identifiability problem [34].

147 Let $\theta_j = (\alpha_j, \mu_j, \sigma_j)$, then $p(x_i)$ has M Gaussian models, and M sets of
148 parameters are needed to be estimated. If $\Theta = (\theta_1, \theta_2, \dots, \theta_M)$, The likelihood function
149 of the GMM model can be expressed as:

$$150 \quad l(x | \Theta) = \log \prod_{i=1}^N \sum_{j=1}^M \alpha_j N_j(x; \mu_j, \sigma_j) = \sum_{i=1}^N \log \sum_{j=1}^M \alpha_j N_j(x; \mu_j, \sigma_j) \quad (3)$$

151 The analytical solution to maximize Equation (3) is generally impractical due to the
152 composite operation of component wise product. The Expectation-Maximization
153 (EM) algorithm is usually applied to generate the unknown parameters (i.e. $\alpha_j, \mu_j, \sigma_j$) in
154 a Gaussian mixture model. The EM algorithm is an iterative procedure for estimating
155 the parameter θ_i of a target distribution that maximize the probability under
156 consideration of a given set of realizations, $\{x_1, x_2, \dots, x_N\}$ [63]. The EM algorithm is
157 an iterative succession of expectation and maximization steps for obtaining the
158 maximum likelihood (ML) estimate, which involves two steps: E-step and M-step. A
159 brief description of the EM algorithm can expressed as follows:

160 E-step: Calculate the posterior probability of mixture component j having generated
161 realization x_i based on the present estimates:

$$162 \quad \beta_{ij} = E(\alpha_j | x_i; \Theta) = \frac{\alpha_j N_j(x_i; \Theta)}{\sum_{j=1}^M \alpha_j N_j(x_i; \Theta)}, \quad 1 \leq i \leq N, \quad 1 \leq j \leq M. \quad (4)$$

163 M-step: Update the model parameters in accordance with their weighted averages
164 across all realizations:

$$165 \quad \alpha'_j = \frac{\sum_{i=1}^N \beta_{ij}}{N} \quad (5)$$

$$166 \quad \mu'_j = \frac{\sum_{i=1}^N \beta_{ij} x_i}{\sum_{i=1}^N \beta_{ij}} \quad (6)$$

$$167 \quad \sigma'_j = \frac{\sum_{i=1}^N \beta_{ij} (x_i - \mu'_j)^2}{\sum_{i=1}^N \beta_{ij}} \quad (7)$$

168

169 **2.2. Copula Method for Bivariate Flood Frequency Analysis**

170 2.2.1. Concept of Copula

171 A copula function is a multivariate probability distribution with its marginal
 172 distribution being uniform. Sklar's Theorem states that any n-dimensional distribution
 173 function F can be formulated through a copula and its marginal distributions, which is
 174 expressed as follows:

$$175 \quad F(x_1, x_2, \dots, x_n) = C(F_{X_1}(x_1), F_{X_2}(x_2), \dots, F_{X_n}(x_n)) \quad (8)$$

176 where $F_{X_1}(x_1), F_{X_2}(x_2), \dots, F_{X_n}(x_n)$ are marginal distributions of random vector $(X_1,$
 177 $X_2, \dots, X_n)$. If these marginal distributions are continuous, then a single copula
 178 function C exists, which can be written as [46, 56]:

$$179 \quad C(u_1, u_2, \dots, u_n) = F(F_{X_1}^{-1}(u_1), F_{X_2}^{-1}(u_2), \dots, F_{X_n}^{-1}(u_n)) \quad (9)$$

180 More details on theoretical background and properties of various copula families can
 181 be found in [46] and [56].

182 A number of copula functions have been developed, mainly including the
 183 Archimedean, elliptical, extreme value copulas. Among them, the Archimedean
 184 copulas are quite attractive in hydrologic frequency analysis, because they can be
 185 easily generated, and are capable of capturing a wide range of dependence structure
 186 with several desirable properties, such as, symmetry and associativity [22]. The Ali-

187 Mikhail-Haq, Cook-Johnson and Gumbel-Hougaard and Frank copulas are most
 188 widely used Archimedean copulas for probabilistic assessment of flood risk. Table 1
 189 presents some basic characteristics of the applied single-parameter bivariate
 190 Archimedean copulas.

191 -----

192 Place Table 1 here

193 -----

194 2.2.2. Conditional Distribution

195 If an appropriate copula function is selected, the conditional joint distribution can
 196 then be obtained. Following [46] and [56], the conditional distribution function of U_1
 197 given $U_2 = u_2$ can be expressed as:

$$198 \quad C_{U_1|U_2=u_2}(u_1) = P(U_1 \leq u_1 | U_2 = u_2) = \frac{\partial}{\partial u_2} C(u_1, u_2) \quad (10)$$

199 Similar conditional cumulative distribution for U_2 given $U_1 = u_1$ can be obtained.
 200 Moreover, the conditional cumulative distribution function of U_1 given $U_2 \leq u_2$ can be
 201 expressed as:

$$202 \quad C_{U_1|U_2 \leq u_2}(u_1) = P(U_1 \leq u_1 | U_2 \leq u_2) = \frac{C(u_1, u_2)}{u_2} \quad (11)$$

203 Likewise, an equivalent formula for the conditional distribution function for U_2 given
 204 $U_1 \leq u_1$ can be obtained.

205 The probability density function (pdf) of a copula function can be expressed as:

$$206 \quad c(u_1, u_2) = \frac{\partial^2 C(u_1, u_2)}{\partial u_1 \partial u_2} \quad (12)$$

207 and the joint pdf of the two random variables can be obtained as:

$$208 \quad f(x_1, x_2) = \frac{\partial^2 C(u_1, u_2)}{\partial x_1 \partial x_2} = \frac{\partial^2 C(u_1, u_2)}{\partial u_1 \partial u_2} \frac{\partial u_1}{\partial x_1} \frac{\partial u_2}{\partial x_2} = f_{x_1}(x_1) f_{x_2}(x_2) c(u_1, u_2) \quad (13)$$

209 Consequently, the conditional pdf of X_1 , given the value of X_2 , can be formulated as:

$$210 \quad f(x_1 | x_2) = \frac{f(x_1, x_2)}{f_{X_2}(x_2)} = f_{X_1}(x_1)c(u_1, u_2) \quad (14)$$

211 And the conditional pdf of X_2 , given the value of X_1 , can be expressed as:

$$212 \quad f(x_2 | x_1) = \frac{f(x_1, x_2)}{f_{X_1}(x_1)} = f_{X_2}(x_2)c(u_1, u_2) \quad (15)$$

213

214 2.2.3. Primary and Secondary Return Period

215 If appropriate copula functions are specified to reflect the joint probabilistic
 216 characteristics among peak, duration and volume of the flood, some conditional,
 217 primary and secondary return periods can be obtained. Specifically, Joint (primary)
 218 return periods called OR and AND can be formulated as [28, 56, 65]:

$$219 \quad T_{u_1, u_2}^{OR} = \frac{\mu}{1 - C_{U_1 U_2}(u_1, u_2)} \quad (16)$$

$$220 \quad T_{u_1, u_2}^{AND} = \frac{\mu}{1 - u_1 - u_2 + C_{U_1 U_2}(u_1, u_2)} \quad (17)$$

221 where μ is the mean inter arrival time of the two consecutive flood events.

222 The secondary return period, called Kendall's return period, is firstly introduced
 223 by Salvadori and De Michele [53] to characterize probability of occurrence of an event
 224 in the area over the copula level curve of value t . This concept has been successively
 225 elaborated and extended by many research works [14, 54-55, 57]. The secondary return
 226 period can be expressed as follows:

$$227 \quad \bar{T}_{u_1, u_2} = \frac{\mu}{1 - K_C(t)} \quad (18)$$

228 where K_C is the Kendall's distribution, associated with theoretical copula function C_θ .

229 For Archimedean copulas, K_C can be expressed as [46]:

$$230 \quad K_c(t) = t - \frac{\phi(t)}{\phi'(t^+)} \quad (19)$$

231 where $\phi'(t^+)$ is the right derivative of the copula generator function $\phi(t)$, as presented in
 232 Table 1.

233

234 2.2.4. Bivariate Hydrologic Risk Analysis

235 Risk is the probability of occurrence of an extreme, dangerous, hazardous, or
 236 (more generally) undesirable event [38]. In engineering design of hydrologic
 237 infrastructures, risk can be explained as the chance of downstream flood attributable
 238 to uncontrolled water release from upstream flood facilities (e.g. a reservoir), leading
 239 to life and property losses [23]. Yen [73] proposed a formulation for the risk of failure
 240 associated with the return period of a flood event, which can be expressed as:

241

$$242 \quad R = 1 - (1 - p)^n = 1 - q^n = 1 - (1 - 1/T)^n \quad (20)$$

243

244 where R is the risk of failure; p and q is the exceedance and nonexceedance
 245 probability, respectively; T is the return period of a flood event; n is the design life of
 246 the hydraulic structure.

247 In practical flood control practice, it is necessary to characterize the flood event
 248 through multiple aspects (e.g. peak and duration) rather than only one flood variable
 249 (e.g. peak). For example, a flood event with high peak flow and long duration may
 250 result in serious losses in properties, while a short-duration event with high peak may
 251 only cause a flash flood. Consequently, bivariate hydrologic risk would be much
 252 helpful in taking nonstructural safety measures, and developing flood mitigation
 253 strategies. In this study, the joint return period in “AND” case is applied to define the

254 bivariate risk analysis as follows:

255

$$256 \quad R_{u_1, u_2} = 1 - \left(1 - \frac{1}{T_{u_1, u_2}^{AND}}\right)^n \quad (21)$$

257

258 **2.3. Goodness-of-fit Statistical Tests**

259 After parameter estimation for both the marginal and joint distributions, the
260 goodness-of-fit statistic tests would be performed to determine whether those
261 estimated distributions are satisfied. The root mean square error (RMSE), Akaikes
262 Information Criterion (AIC) and the Kolmogorov-Smirnov (K-S) goodness-of-fit tests
263 would be employed to evaluate the performance of the marginal distributions obtained
264 through the parametric distributions and the Gaussian mixture model (GMM). And the
265 Rosenblatt transformation [49] would be applied to investigate the performance of
266 joint distributions in describing the dependency between flood variable pairs.

267 In the process of evaluating the performance of marginal distributions obtained
268 through the parametric methods and GMM, the empirical nonexceedance probabilities
269 would be obtained through the Gringorten plotting position formula [30], which is
270 expressed as:

$$271 \quad P(K \leq k) = \frac{k - 0.44}{N + 0.12} \quad (22)$$

272 where N stands for the sample size; k stands for the k^{th} smallest observation in the
273 data set; and the data set is arranged in an increasing order.

274 The RMSE, Akaikes Information Criterion (AIC) and the K-S test are used to
275 evaluate fitting effect of different probability distributions to the flood variables.

276 The RMSE can be expressed as [69]:

$$277 \quad RMSE = \sqrt{\frac{\sum_{k=1}^N (x_k^{est} - x_k^{obs})^2}{N}} \quad (23)$$

278 where x_k^{est} denote theoretical values from the fitted probability distribution; x_k^{obs} denote
 279 the empirical probabilities obtained through Equation (22); N is the sample size.

280 Based on RMSE, the AIC value can be obtained as follows:

$$281 \quad AIC = N * \ln((RMSE)^2) + 2k \quad (24)$$

282 where k is the number of unknown parameters in the probability distribution.

283 The K-S test is a nonparametric probability distribution free test [80]. The
 284 statistic of K-S test quantifies the largest vertical difference between the estimated and
 285 empirical distributions [44, 47]. Given n increasing ordered data points, $x_{(.)}$, the K-S
 286 test statistic is defined as [11]:

$$287 \quad T = \sup_x |F^*(x) - F_n(x)| \quad (25)$$

288 where $F^*(x)$ means the estimated distribution, $F_n(x)$ denotes the empirical
 289 distribution, and ‘sup’ stands for supremum. The P-value for K-S test was
 290 approximated using Miller’s approximation [80].

291 For evaluating the performance of copulas, the goodness-of-fit test based on
 292 Rosenblatt transformation would be employed based on the recommendation of
 293 Genest et al. [25]. They argued that test statistics based on the Cramér von Mises
 294 functional of a process tend to be more powerful than those based on the
 295 Kolmogorov–Smirnov distance taken on the same process [25]. Consequently,
 296 Cramér von Mises statistic will be adopted to test the performance of the copulas with
 297 the corresponding p-values being approximated through Monte Carlo simulation. The
 298 detailed procedures for performing goodness-of-fit test for copulas based on
 299 Rosenblatt transformation are provided by Genest et al. [25].

300

301 **3. Study Area and Data**

302 **3.1. Overview of the Studied Watershed**

303 The proposed GMM-copula method would be applied to the Yangtze River to
304 demonstrate the applicability of the proposed method in analyzing multivariate flood
305 risk. Yangtze River is the longest river in Asia, and the third longest river in the world,
306 with a length of 6,300 km for the main stream, flowing from Qinghai Province
307 eastward to the East China Sea at Shanghai. Floods of the Yangtze River in central
308 and eastern China have occurred periodically and often caused considerable
309 destruction of property and loss of life [10]. For example, in 1998, the entire Yangtze
310 River basin suffered from tremendous flood—the largest flood since 1954, which led
311 to the economic loss of 166 billion Chinese Yuan [74]. Hence, multivariate flood risk
312 analysis for Yangtze River is very important for flood prevention and disaster relief.

313 For the Yangtze River, floods are caused by temporal-spatial variation in
314 precipitation. A large part of the Yangtze River Basin has subtropical monsoon
315 climate, with the precipitation being concentrated during summer season.
316 Consequently, summer is the main flood season due to the heavy monsoon rainfall
317 [10]. The floods in the middle and lower reaches of the Yangtze River mainly stem
318 from the upper region of the Yichang Station. The Yichang station plays a vital role
319 for flood control in the middle and lower reaches of Yangtze River. It is also the
320 control station for the Three Georges Reservoir. The flood from Yichang station
321 contributes about 50% of the total flow volume of the middle and lower reaches of
322 Yangtze River. Moreover, The Jingjiang reach (Figure 1), located in the middle reach
323 of the Yangtze River from Zhicheng to Chenglingji with a length of 340 km, is the

324 most prone area to suffer floods in the Yangtze River basin. Approximate 90% flood
325 in Jingjiang reach comes from the flood in Yichang station [9].

326 Due to the key role of the Yichang station in controlling the flood in the middle
327 and lower reaches of Yangtze River, the daily streamflow data from Yichang station
328 would be applied to analyze the bivariate flood risk in Yangtze River. Figure 1 shows
329 the location of Yichang station, which is also the control site of the Three Gorges Dam
330 (TGD). The Three Gorges Dam (TGM) is the largest hydraulic project in terms of
331 design capacity over the world. It has produced dramatic benefits in flood control,
332 power generation and navigation. Recently, the impacts of the TGM project on
333 hydrology and environment have been attracting the world's attention. The Yichang
334 Station is the control site of TGD, which also divided the Yangtze River into the upper
335 and middle reaches. This study mainly focused on the flood from the upper Yangtze
336 River, which is 4,529 km long, up to 3/4 of the whole length of the Yangtze River,
337 with a drainage area of 1,006,000 km² [8].

338

339 **3.2. Historical Flood Characteristics at Yichang station**

340 Based on the daily flow data, the annual maximum peak discharges and the
341 corresponding hydrograph volumes and durations values can be obtained. Hence,
342 although the peak discharges are definitely annual maximums, the hydrograph
343 volumes and durations are not necessarily also annual maximums [65]. The single-
344 peaked flood hydrograph is shown in Figure 2. Flood duration (D) can be determined
345 by identifying the time of rise (point “s” in Figure 2) and fall (point “e” in Figure 2) of
346 the flood hydrograph. The start of the surface runoff is marked by the sharp rise of the
347 hydrograph and end of the flood runoff is identified by the inflection point on the
348 receding limb of the hydrograph. Between these two points, the total flood volume is

349 estimated. If time of rise of the flood hydrograph is denoted by SD (day) and fall by
 350 ED (day), the flood volume (V) of each flood event is determined using following
 351 expression (Yue 2001):

$$352 \quad V_i = (V_i^{total} - V_i^{baseflow}) = \sum_{j=SD_i}^{ED_i} Q_{ij} - \frac{1}{2}(Q_{is} + Q_{ie})(1 + D_i) \quad (29)$$

353 For a flood with multiple peaks, the peak flow would be the maximum peak value in
 354 the flood. The corresponding duration is identified based on Figure 2 and the
 355 associated volume is calculated through Equation (29). Moreover, when multiple
 356 floods happen in one year, the flood with maximum peak is only considered since risk
 357 analysis pays attention to flood extremes. Once the flood characteristics are obtained
 358 from daily streamflow data, then flood frequency analysis can be analyzed. Figure 3
 359 shows the variations in flood peak discharge (i.e., Q (m³/s)), hydrograph volume, (i.e.,
 360 V (m³/s day)) and hydrograph duration, (i.e., D (day)) from 1882-2007.

361

362 -----

363 Place Figures 2 and 3 here

364 -----

365

366 **4. Result Analysis**

367 **4.1. Marginal Probability Distribution Functions of Flood Variables**

368 One of the main advantages for the copula method is that the marginal
 369 distributions and multivariate dependence modelling are two separate processes.
 370 Consequently, to analyze the multivariate flood frequency in the Yangtze River, the
 371 marginal distributions of flood variables can be quantified firstly. In this study, the
 372 Gaussian mixture model would be applied to quantify the marginal distributions of

373 flood peak, volume and duration. Besides, many parametric distributions have been
374 used to estimate flood frequencies from observed annual flood series, such as the
375 general extreme value distribution in the United Kingdom, Log-Pearson Type-III in
376 the U.S. and Pearson Type III in China [1, 37, 70, 75]. To demonstrate the
377 performance of GMM in modeling the marginal distributions of flood variables, the
378 GMM would be compared with four parametric methods, including Gamma, GEV,
379 Lognormal distributions and Pearson Type III. The expressions for probability
380 functions (pdfs) for Gamma, GEV, Lognormal, Pearson Type III and the values of
381 their associated unknown parameter are presented in Table 2. These parameters are
382 obtained through maximum likelihood estimation method. Table 3 shows the marginal
383 distributions of flood variables obtained through GMM, in which the unknown
384 parameters are obtained through the EM algorithm.

385

386 -----

387 Place Tables 2 and 3 here

388 -----

389

390 Figure 4 illustrates the fitted marginal distributions for the three flood variables
391 through Gamma, GEV, Lognormal, Pearson Type III (i.e. P3), and GMM-based
392 distribution functions. The cdfs and pdfs for the marginal distributions of flood
393 variables (in Figure 3) show good agreement between the theoretical and the
394 empirical distributions. Generally, the flood peak and volume can be well quantified
395 through the proposed four parametric distributions and GMM-based distributions. For
396 the flood duration, there are some deviations between the theoretical and observed
397 values, especially for the four parametric distributions. To further evaluate the GMM

398 and four parametric distributions in quantifying the probability distributions of flood
399 variables, the Kolmogorov-Smirnov (K-S) test would be conducted. Table 4 presents
400 the results of K-S tests. The results indicate that all the proposed five methods can be
401 employed to model the distributions of flood peak, volume and duration, with the P-
402 values larger than 0.05. However, the performance of the four parametric distributions
403 in modelling the flood duration is not as well as those in quantifying the flood peak
404 and volume, since the P-values are less than 0.1. The root mean square error (RMSE)
405 and AIC values, which are respectively expressed as Equations (23) and (24), would
406 then adopted to compare the performance of those four distributions. As shown in
407 Table 4, the GMM-based distributions perform best in quantifying the three flood
408 variables, with lowest RMSE and AIC values. Especially for flood duration, the
409 GMM-based distribution performs much better than the other four parametric
410 distributions.

411

412 -----

413 Place Figure 4 and Table 4 here

414 -----

415

416 **4.2. Joint Distributions Based on Copula Method**

417 The dependence of flood variables was evaluated through the Pearson's linear
418 correlation (r), and one non-parametric dependence measure, Kendall's tau. Table 5
419 presents the values of Pearson's linear correlation coefficient and Kendall's tau
420 among flood peak, volume and duration. The values of Pearson's r and Kendall's tau
421 between duration and volume are highest, followed by the flood pairs of peak-volume,
422 and peak-duration. In detail, the Pearson, Kendall correlation coefficient values are

423 0.55 and 0.66 for peak-volume, 0.68 and 0.75 for volume-duration, and 0.27 and 0.35
424 for peak-duration. These results indicate that the correlation between the flood
425 duration and volume would be higher than the other two flood variable pairs. In our
426 case, the correlation coefficient for peak and duration is much smaller than for the
427 other two pairs (i.e. peak-volume and volume-duration), which is consistent with
428 conclusions from previous studies [29, 36, 48, 65].

429 The Archimedean copulas are the most attractive copulas for multivariate
430 hydrologic risk analysis due to their ease for construction and capability of capturing
431 dependence structure with several desirable properties. The Cook-Johnson (Clayton),
432 Gumbel-Hougaard, Frank and Ali-Mikhail-Haq copulas are the four widely used
433 Archimedean copulas. However, the Ali-Mikhail-Haq copula is only applicable with
434 the Kendall's tau value varied within $[-0.18, 0.33]$ [46]. In this study, the flood pair of
435 peak-duration exhibits the lowest Kendall's tau values with a value being 0.35. For
436 the flood pairs of peak-volume and volume-duration, the corresponding Kendall's tau
437 values are 0.66 and 0.78, respectively. Consequently, the Ali-Mikhail-Haq copula is
438 excluded and the Cook-Johnson (Clayton), Gumbel-Hougaard and Frank copulas
439 would be selected to model the dependence among flood variables. The unknown
440 parameters in these four copulas are estimated by method-of-moments-like (MOM)
441 estimator based on inversion of Kendall's tau.

442 The joint distribution functions for flood peak and volume, obtained through the
443 three above-mentioned copulas, are shown in Figure 5; the joint distributions for
444 peak-duration, and volume-duration are shown in Figures 6 and 7, respectively. Also,
445 comparison between empirical and theoretical copula functions for the flood pairs of
446 peak-volume, peak-duration and volume-duration can be found in Figure 5, 6 and 7,

447 respectively. In Figures 5 to 7, the red dashed contour lines represent the empirical
448 copula obtained through $C_n(u, v) = 1/n \sum_{i=1}^n 1(R_i / (n+1) \leq u, S_i / (n+1) \leq v)$, where
449 $u, v \in [0, 1]$, R_i and S_i denote the ranks of the ordered sample, and the solid contour
450 lines represent the theoretical copula. The results indicate that the empirical and the
451 three theoretical copulas can match well for flood peak-volume. For flood peak-
452 duration, there are some deviations between theoretical copulas and empirical copula
453 at low probability levels. This may due to the discrete characteristic of the duration
454 sample and the relative low accuracy of the obtained marginal distribution. However,
455 at high probability levels, the theoretical values can fit well with the empirical copula
456 values. Also, similar characteristic can be found for the flood pair of volume-duration.

457 Since there are three candidate copulas, investigating the differences among the
458 three chosen copulas and identifying the most appropriate copulas for further analysis
459 are necessary. In this study, the Rosenblatt transformation with Cramér von Mises
460 statistic is employed to evaluate performance of the proposed three copulas in
461 modelling joint distributions of flood variable pairs. Table 6 presents the results of
462 statistic test results for the three flood pairs. It can be seen that the proposed Cook-
463 Johnson (Clayton), Gumbel-Hougaard and Frank copulas can be applicable for
464 modelling the dependence of flood peak-volume, peak-duration and volume-duration,
465 with the p-values larger than 0.05. To further identify the most appropriate one, the
466 root mean square error (RMSE) (expressed by Equation (23)) is used to test the
467 goodness of fit of sample data for the theoretical joint distribution obtained using
468 copula functions. Table 6 shows the RMSE values for joint distributions obtained
469 through different copula functions for flood peak-volume, peak-duration and volume-
470 duration. The differences among these three copulas in quantifying the joint
471 probabilities of the three flood pairs are rarely small. Take the flood pair of peak-

472 volume as an example, the RMSE value for the Gumbel-Hougaard and Cook-Johnson
473 copula is 0.0168 and 0.0199 respectively, while the RMSE value of Frank copula is
474 0.0149. Based on the values of RMSE, it can be concluded that the Frank copula
475 would be best for quantifying the joint distribution of flood peak-volume. Similarly,
476 the Frank copula would be the most appropriate copula for modelling the joint
477 distribution of flood peak-duration and volume-duration.

478

479

480 -----

481 Place Figures 5 - 7 here and Table 6

482 -----

483 **4.3. Bivariate and Conditional Risk Analysis**

484 4.3.1. Conditional Cumulative Distribution Functions and Return Periods of Flood 485 Characteristics

486 Based on the results presented in Table 6, the Frank copula would be chosen to
487 model the dependence between the three flood pairs. Consequently, the conditional
488 cumulative distribution functions (cdfs) of one flood variable, given the value of the
489 other flood variable value, can be derived based on the fitted copula function.

490 Figure 8 shows the conditional cdfs of flood variables, which are obtained through
491 Equations (10) and (11). It can be seen that, among the flood pairs of peak-volume,
492 peak-duration and volume-duration, the values of conditional cdf for one flood
493 variable would decrease as the value of other flood variable increase. This indicates
494 positive correlation structures between peak-volume, peak-duration, and volume-
495 duration. Besides, the decreasing trend of conditional cdfs for peak-duration is less
496 than the other two pairs, indicating less correlation structures between peak and

497 duration. This is consistent with the results presented in Table 5.

498

499 -----

500 Place Figure 8 Here

501 -----

502

503 The concurrence probabilities of various combination of flood variable would be
504 more helpful for actual flood control and management than the univariate flood
505 frequency analysis. As expressed as Equations (16) – (19), the joint return period and
506 second return period can be derived based on the selected copula functions. Table 7
507 presents the joint return periods of “AND” and “OR” cases for different flood pairs.
508 In general, the joint return period in “AND” case is much longer than the joint return
509 period in “OR” case. For example, if both the flood peak and duration are in 100-year
510 return period, the “OR” joint return period of flood peak-duration would be 50.9
511 years, while, in contrast, the “AND” joint return period is 2809.4. Furthermore, the
512 “AND” return period for flood peak-duration is longest among the three flood
513 variable pairs due to the low correlation between flood peak and duration, followed by
514 the “AND” return periods of peak-volume and volume-duration. Correspondingly, the
515 “OR” joint return period of peak-duration is shorter than the “OR” return periods of
516 the other two flood variable pairs. Figure 9 shows the contour plot of the joint return
517 periods in “OR” and “AND” cases for different flood pairs. Also, the secondary return
518 periods are presented in Table 7, which can be useful for analyzing risk of
519 supercritical flood events. The secondary return period is defined as the average time
520 between the concurrence of two supercritical flood events, which would appear more
521 rarely than the given design return period. As the primary return period increases, the

522 probability of supercritical flood events decreases, leading to increase of the
523 secondary return period. Furthermore, the secondary return period is higher than the
524 joint return period in T^{OR} case but less than the joint return period in T^{AND} case.

525 -----

526 Place Table 7 and Figure 9 here

527 -----

528

529 4.3.2. Bivariate Hydrologic Risk Analysis

530 The damages caused by a flood, such as the failure of hydraulic structures,
531 mainly due to the high peak flow of the flood. The annual maximum peak discharge
532 would be the central issue to be considered for hydrologic risk analysis. Moreover, the
533 flood discharge volume and duration would be also under consideration in practical
534 flood control and mitigation, in which the flood duration is the vital factor for
535 decision maker in characterizing the flood control pressure, and the flood volume is
536 related to flood diversion practices. Consequently, multivariate flood risk analysis,
537 which involves more flood variables than just considering flood peak, would be more
538 helpful for actual flood control. Therefore, in this study, a bivariate hydrologic risk
539 analysis method would be proposed to identify the inherent flood characteristics in
540 Yangtze River. In particular, three flow amounts, with a return period of 50, 70, and
541 100-year, respectively are considered as designed standard for the river levee around
542 the Yichang Station. Four service time scenarios are also assumed for the river levee,
543 namely 30, 50, 70 and 100 years.

544

545 (1) Bivariate flood risk under different flood peak-volume scenarios

546 The bivariate hydrologic risk for flood peak flow and volume indicates the

547 concurrence probabilities of flood peak flow and volume values. Figure 10 shows the
548 bivariate flood risk under different flood peak-volume scenarios. For the univariate
549 hydrologic risk expressed as Equation (20), its value would decrease as the increase of
550 designed peak flow or the service time of the river levee. As can be seen from Figure
551 10, if the flood volume is less than 1×10^5 ($\text{m}^3/\text{s day}$), the bivariate risk values for
552 flood peak-volume would not decrease significantly for all designed flows and service
553 time periods. This suggests that the occurrence of one flood peak flow would usually
554 be accompanied with a flood volume up to 1×10^5 ($\text{m}^3/\text{s day}$). However, for one
555 designed flow and service time period, the values of the bivariate risk for flood peak-
556 volume would decrease when the associated flood volume is larger than 1×10^5 (m^3/s
557 day). This indicates that the probabilities of concurrence of large flood volumes and
558 high peak flows would be generally less than the occurrence probabilities of high
559 peak flows.

560

561 -----

562 Place Figure 10 here

563 -----

564

565 The implication for the bivariate risk of flood peak flow and volume is to provide
566 decision support for hydrologic facility design and establishment of flood diversion
567 areas. In actual flood control practices, the excess water of floods can be redirected
568 temporary holding ponds or other bodies of water with a lower risk or impact to flood.
569 For example, in China, the flood diversion areas are rural areas that are deliberately
570 flooded in emergencies in order to protect cities. In flood diversion practice, the
571 bivariate risk for flood peak flow and volume would be an important reference for the

572 design of flood diversion areas. For example, as shown in Figure 9, for the river levee
573 with a designed flow of 50-year return period and 30-year service period, the flood
574 risk value would be about 45, 43, 35, 22% with a flood volume being 0.5, 1, 1.5, and
575 $2 \times 10^5 \text{ m}^3/\text{s}$, respectively. Based on these bivariate risk values, the flood manager can
576 design corresponding scales of the flood diversion areas.

577

578 (2) Bivariate flood risk under different flood peak-duration scenarios

579 Figure 11 shows the variations in the failure risk of river levee around Yichang
580 Station under different flood peak-duration scenarios. The bivariate hydrologic risk
581 can reflect the failure risk of river levee with respect to the variation of flood
582 durations. In Figure 11, the initial risk values (points on the y-coordinate) are obtained
583 through Equation (19) without considering impacts of the flood duration, while the
584 points on the solid, dashed and asterisk lines are derived based on Equation (21). The
585 results in Figure 10 indicate that the bivariate risk of flood peak-duration would not
586 decrease at the flood duration less than 8 days, and then decrease as the increase of
587 flood duration. Such results suggest that the once a flood occurs at Yichang Station,
588 this flood would last up to 8 days without significant decrease in the occurrence
589 probability. However, the concurrence of a flood with high peak flow and long
590 duration would not appear frequently.

591

592 -----

593 Place Figure 11

594 -----

595

596 The bivariate risk of the flood peak flow and duration can provide useful

597 information for actual hydrologic facility design and potential flood control. In
598 practical engineering hydrologic facility construction, the return period of peak flow
599 would be the key factor to be considered. Moreover, the flood duration would be
600 related to flood defense preparation, in which longer flood duration would generally
601 require more flood defense materials such as sand, wood, bags. Consequently, the
602 bivariate flood risk values under different flood peak-duration scenarios would be
603 considered as references for decision makers to determine how much materials would
604 be prepared for flood defense.

605

606 4.3.3. Conditional Probability Density Functions of Flood Characteristics

607 In addition to derive the conditional cdfs and joint return periods based on the
608 best-fitted copula for the historical flood data, the conditional probability density
609 functions (pdfs) of the flood variable can also be generated based on Equations (12) -
610 (15). In flood risk analysis, the peak flow would be the critical factor to judge whether
611 a flood appears. However, once the flood occurred, the severity of the flood would
612 also influenced by flood duration and volumes. In detail, the flood duration would be
613 related to the flood control pressure in which flood defense materials should be
614 prepared for strengthening the river levee and inspection should be conducted for the
615 safety of the river levee. The flood volume would generally influence the flood
616 diversion practices, in which excess water would be diverted to temporary holding
617 ponds with lower risk in order to protect cities.

618 Figure 12 shows the distributions of flood volume conditional on the flood peak
619 flows with different return periods. In this study, the flood peak flows with return
620 periods of 10, 20, 50, and 100-year are under consideration. Each curve represents the
621 probability distribution function (pdf) of flood volume associated with the flood peak

622 flow with a particular return period. It can be seen that, once a flood appears, the
623 conditional pdf of flood volume would approximately follow a bimodal distribution,
624 with the two vertexes appearing around 1.2 and 2.0×10^5 m³/s day, respectively. More
625 specifically, the former vertex would appear more frequently for small floods while
626 the latter one is more frequent for large floods. Moreover, as the peak flows increases,
627 the two vertexes of the flood volume would also increase correspondingly, but the
628 latter vertex seems to increase more than the former vertex, as shown in Figure 12.
629 Finally, the conditional pdf of flood volume also shows that the occurrence
630 probability of the first vertex would decrease while the occurrence probability of the
631 latter vertex would increase when the return period of the flood peak increases. Such
632 pdfs of flood volume conditional on different flood peak flows can provide support
633 information for flood diversion practices and be involved in the flood optimization
634 models to determine the capacities of flood diversion. For instance, once a flood
635 occurs and excessive flood is required to be diverted to some flood discharge area, the
636 associated flood volume should be estimated before conducting flood diversion. From
637 Figure 12, it can be concluded that two flood volumes would be primarily under
638 consideration, around 1.2 and 2.0×10^5 m³/s day, respectively. Particularly, the flood
639 volume of 1.2×10^5 m³/s day would be paid more attention for small floods while the
640 volume of 2.0×10^5 m³/s day would be paid more attention for large floods. These
641 results can provide useful information for flood managers to prepare appropriate flood
642 diversion schemes. Moreover, Table 8 shows the statistical characteristics for the
643 PDFs of flood volume conditional on different floods. The results indicate that, as the
644 increase of the flood peak return period, the mean value of the conditional pdf of
645 flood volume would generally increase, while the standard deviation of the
646 conditional pdfs would not change significantly.

647

648 -----

649 Place Figure 12 and Table 8 here

650 -----

651

652 Figure 13 shows the distributions of flood duration conditional on the flood peak
653 flows with different return periods. It is indicated that, the conditional pdfs of flood
654 duration would also obey bimodal distributions, with two vertexes appearing around
655 11 and 15 days. Specifically, as the increase of flood return period, the former vertex
656 around 11 days would not change significantly, but the latter vertex (around 15 days)
657 would show a remarkable increase. For instance, the latter vertex of the duration pdf
658 conditional on a flood with 10 years would be about 14.8 days, while such a vertex
659 would increase to around 15.7 days when the return period of the flood peak increase
660 to 100 years. Moreover, the latter vertex show a more frequent occurrence probability
661 than the former vertex except for a small flood with a 10-year return period. The
662 engineering implications of the pdfs of flood duration conditional on flood peak is to
663 provide an insightful screening for the duration time once a flood occurs, which will
664 further be considered as a reference for flood defense materials preparation and river
665 levee safety inspection. The statistical characteristics of the conditional pdfs are
666 presented in Table 8. The results indicate that the mean values of the conditional pdfs
667 would increase while the standard deviations are nearly constants. Furthermore, for a
668 flood with a return period larger than 50 years, associated mean value of the flood
669 duration would not change significantly, even though the latter vertex shows apparent
670 increase with the increase in flood peak return period.

671

672 -----

673 Place Figure 13 here

674 -----

675

676 **5. Conclusions**

677 In this study, a bivariate hydrologic risk analysis method is proposed through coupling
678 Gaussian mixtures into copulas. In the bivariate hydrologic risk analysis framework,
679 the bivariate frequency analysis, which considered the flood variables pairs of flood
680 peak, duration and volume, was firstly conducted through coupling Gaussian mixture
681 models into copulas, leading to a coupled GMM-copula method. This method
682 improved upon previous methods through providing better estimation for marginal
683 distributions through Gaussian mixture models. The primary, conditional and
684 secondary return periods were then derived based on the selected copula. The
685 bivariate hydrologic risk was defined based on the joint return period of flood
686 variables to reflect the hydrologic risks of flood peak-duration and flood peak-volume
687 pairs. Besides, the conditional probability distribution functions (pdfs) of flood
688 volume and duration under different flood peak scenarios were also derived to explore
689 the variation in pdfs of flood volume and duration corresponding to different flood
690 peak flows.

691 The proposed method was applied for quantifying the bivariate hydrologic risk in
692 the Yangtze River based on the daily discharge measurements at Yichang Station. The
693 results indicated that, compared with the parametric distributions such as Gamma,
694 GEV and Lognormal and Pearson Type III functions, the Gaussian mixture model
695 could perform much better for quantifying the marginal distributions of flood peak,
696 volume and duration. Such conclusions has been demonstrated through the K-S test,

697 the RMSE and AIC values.

698 For the dependence among flood variables, the Frank copula would be best for
699 quantifying the joint distributions of the three flood variable pairs. The bivariate risks
700 of flood peak-volume and flood peak-duration were evaluated based on the joint
701 return period in “AND”, revealing significance of effects from persisting high risk
702 levels due to impacts from multiple interactive flood variables. The results show that
703 the bivariate risk of flood peak-volume would keep constant for the corresponding
704 volume less than $1.0 \times 10^5 \text{ m}^3/\text{s}$ day, show apparent decrease for the flood volume
705 varying between 1.0 and $1.7 \times 10^5 \text{ m}^3/\text{s}$ day, and present most significant decreasing
706 rates for the volume larger than about $1.7 \times 10^5 \text{ m}^3/\text{s}$ day. For the bivariate risk of flood
707 peak-duration, it would not change significantly for the flood duration less than about
708 8 days and then show significant decreasing rate. Moreover, the pdfs of flood volume
709 and duration conditional on flood peak appeared to be bimodal. The two vertexes for
710 the conditional pdfs of flood volume were located at around 1.2 and $2.0 \times 10^5 \text{ m}^3/\text{s}$
711 day; the occurrence probability for the former vertex would decrease and that for the
712 latter one would increase with the return period of the flood peak increases. The two
713 vertexes for the conditional pdfs of flood duration appeared at around 11 and 15 days,
714 respectively, with the associated occurrence probabilities respectively decreasing and
715 increasing with the increase of the flood peak return period.

716 In engineering applications, the bivariate risk can be applied for actual flood
717 management. Specifically, the bivariate risk of flood peak-volume can provide
718 support for design of flood diversion area, and the bivariate risk of flood peak-
719 duration can be considered as a reference for preparation of flood defense materials.
720 Moreover, the pdfs of flood volume and duration conditional on different flood flows
721 can help flood mitigation and control once a flood has occurred, in which the

722 conditional pdfs of flood volume can provide useful information for flood diversion,
723 and the conditional pdfs of flood duration can help decision maker arrange related
724 people for river levee inspection.

725

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730

731 **References**

- 732 [1] Adamowski K. (1989). A Monte Carlo comparison of parametric and nonparametric
733 estimation of flood frequencies. *Journal of Hydrology*, 108, 295-308
- 734 [2] AghaKouchak, A., A. Bárdossy, and E. Habib (2010a), Conditional simulation of remotely
735 sensed rainfall data using a non-gaussian v- transformed copula, *Advances in Water Resources*,
736 33 (6), 624 – 634, doi:DOI:10.1016/j.advwatres.2010.02.010.
- 737 [3] AghaKouchak, A., A. Bárdossy, and E. Habib (2010b), Copula-based uncertainty modelling:
738 application to multisensor precipitation estimates, *Hydrological Processes*, 24, 2111–2124,
739 doi:10.1002/hyp.7632.
- 740 [4] Aghakouchak, A., G. Ciach, and E. Habib (2010c), Estimation of tail dependence coefficient
741 in rainfall accumulation fields, *Advances in Water Resources*, 33(9), 1142-1149,
742 doi:10.1016/j.advwatres.2010.07.003.
- 743 [5] Bobee B, Cavidas G, Ashkar F, Bernier J, Rasmussen P (1993). Towards a systematic approach
744 to comparing distributions used in flood frequency analysis. *Journal of Hydrology* 142:121–
745 136
- 746 [6] Chebana, F., S. Dabo-Niang, and T. B. M. J. Ouarda (2012), Exploratory functional flood
747 frequency analysis and outlier detection, *Water Resources Research*, 48, W04514.
- 748 [7] Chebana, F., and T. B. M. J. Ouarda (2009), Index flood-based multivariate regional
749 frequency analysis, *Water Resources Research*, 45, doi:10.1029/2008WR007490.
- 750 [8] Chen L., Singh V.P., Guo S., (2013). Measure of correlation between river flows using the
751 copula-entropy method. *Journal of Hydrologic Engineering*, 18, 1591-1606.
- 752 [9] Chen, L., Ye, L., Singh, V., Zhou, J., and Guo, S. (2012). Flood coincidence risk analysis
753 using multivariate copula functions. *Journal of Hydrologic Engineering*, 17(6), 742-755
- 754 [10] Chen, L., Ye, L., Singh, V., Zhou, J., and Guo, S. (2014). Determination of Input for Artificial
755 Neural Networks for Flood Forecasting Using the Copula Entropy Method. *Journal of*

- 756 Hydrologic Engineering, doi,10.1061/(ASCE)HE.1943-5584.0000932,
- 757 [11] Conover W.J. (1999). *Practical nonparametric statistics*. Third Edition, John Wiley & Sons,
- 758 Inc. New York, pp.428-433
- 759 [12] Cunnane C (1989). Statistical distributions for flood frequency analysis. WMO No. 718,
- 760 WMP, Geneva
- 761 [13] De Michele C, Salvadori G (2003) A Generalized Pareto intensity-duration model of storm
- 762 rainfall exploiting 2-copulas. *Journal of Geophysical Research*, 108(D2), 4067,
- 763 doi:10.1029/2002JD002534.
- 764 [14] De Michele C., G. Salvadori, R. Vezzoli, and S. Pecora. Multivariate assessment of droughts:
- 765 frequency analysis and Dynamic Return Period. *Water Resources Research*, 49(10):6985–
- 766 6994, 2013. doi: 10.1002/wrcr.20551.
- 767 [15] Diehl T., Potter K. W. Mixed flood distribution in Wisconsin. *Hydrologic Frequency*
- 768 *Modelling*. Netherlands: D. Reidel Publishing Company, 1987, 213-226.
- 769 [16] Fan Y.R., Huang G.H., Huang K., Baetz B.W., (2015a). Planning Water Resources Allocation
- 770 under Multiple Uncertainties through A Generalized Fuzzy Two-Stage Stochastic
- 771 Programming Method. *IEEE Transactions on Fuzzy Systems*, doi
- 772 10.1109/TFUZZ.2014.2362550
- 773 [17] Fan Y.R., Huang G.H., Li Y.P., (2012). Robust interval linear programming for environmental
- 774 decision making under uncertainty. *Engineering Optimization* 44 (11), 1321-1336
- 775 [18] Fan Y.R., Huang W.W., Huang G.H., Huang K., Li Y.P., Kong X.M., (2015b). Bivariate
- 776 hydrologic risk analysis based on a coupled entropy-copula method for the Xiangxi River in
- 777 the Three Gorges Reservoir area, China. *Theoretical and Applied Climatology*, doi:
- 778 10.1007/s00704-015-1505-z
- 779 [19] Fan Y.R., Huang W.W., Huang G.H., Huang K., Zhou X., (2015c). A PCM-based stochastic
- 780 hydrological model for uncertainty quantification in watershed systems. *Stochastic*
- 781 *Environmental Research and Risk Assessment*, 29, 915-927
- 782 [20] Fan Y.R., Huang W.W., Li Y.P., Huang G.H., Huang K., (2015d). A coupled ensemble
- 783 filtering and probabilistic collocation approach for uncertainty quantification of hydrological
- 784 models. *Journal of Hydrology*, doi: 10.1016/j.jhydrol.2015.09.035
- 785 [21] Feng P., Li X. Bivariate frequency analysis of non-stationary flood time series based on
- 786 Copula methods. *Journal of Hydraulic Engineering*, 2013, 44(10), 1137-1147,
- 787 doi:10.13243/j.cnki.slxb.2013.10.005 (in Chinese with English abstract).
- 788 [22] Ganguli P., Reddy M.J., (2013). Probabilistic assessment of flood risks using trivariate
- 789 copulas. *Theoretical and Applied Climatology*, 111(1-2), 341-360.
- 790 [23] Gebregiorgis A.S., Hossain F., (2012). Hydrological risk assessment of old dams: case study
- 791 on Wilson dam of Tennessee River basin, *Journal of Hydrologic Engineering*, 17, 201-212.
- 792 [24] Genest C, Favre AC. (2007). Everything you always wanted to know about copula modeling
- 793 but were afraid to ask. *Journal of Hydrologic Engineering*, 12(4): 347–368.

- 794 [25] Genest C., Rémillard B., Beaudoin D., (2009). Goodness-of-fit tests for copulas: A review and
795 a power study. *Insurance: Mathematics and Economics* 44:199-213.
- 796 [26] Ghizzoni, T., G. Roth, and R. Rudari (2010), Multivariate skew-t approach to the design of
797 accumulation risk scenarios for the flooding hazard, *Advances in Water Resources*, 33 (10),
798 1187-1290, doi:10.1016/j.advwatres.2010.08.003.
- 799 [27] Ghizzoni T., G. Roth and R. Rudari (2012), Multisite flooding hazard assessment in the
800 Upper Mississippi River, *Journal of Hydrology*, 412-413, 101-113
- 801 [28] Graler B., van den Berg M.J., Vandenberghe S., Petroselli A., Grimaldi S., De Baets B.,
802 Verhoest N.E.C., (2013). Multivariate return periods in hydrology: a critical and practical
803 review focusing on synthetic design hydrograph estimation. *Hydrology and Earth System
804 Sciences* 17: 1281–1296.
- 805 [29] Grimaldi S., Serinaldi F., (2006). Asymmetric copula in multivariate flood frequency
806 analysis. *Advances in Water Resources*, 29, 1281-1296.
- 807 [30] Gringorten I.I. (1963). A plotting rule for extreme probability paper. *Journal Geophysical
808 Research* 68:813-814.
- 809 [31] Han J.C., Huang G.H., Zhang H., Li Z., Li Y.P., (2014). Bayesian uncertainty analysis in
810 hydrological modeling associated with watershed subdivision level: a case study of SLURP
811 model applied to the Xiangxi River watershed, China. *Stochastic Environmental Research
812 and Risk Assessment*, 28(4), 973-989.
- 813 [32] Hao Z. and V.-P. Singh (2012), Entropy-copula method for single-site monthly streamflow
814 simulation, *Water Resources Research*, 48, W06604, DOI:10.1029/2011WR011419.
- 815 [33] Hastie T., Tibshirani R., Friedman J., (2009). *The Elements of Statistical Learning: Data
816 Mining, Inference, and Prediction*, Springer, New York, 2009
- 817 [34] He J., Mixture model based multivariate statistical analysis of multiply censored environmental
818 data. *Advances in Water Resources*, 59, 15-24.
- 819 [35] Kao S.C., Govindaraju R.S., (2010). A copula-based joint deficit index for droughts. *Journal
820 of Hydrology*, 380, 121-134.
- 821 [36] Karmakar S., Simonovic S.P., (2009). Bivariate flood frequency analysis. Part 2: a copula-
822 based approach with mixed marginal distributions. *Journal of Flood Risk Management*, 2,
823 32-44.
- 824 [37] Kidson R., Richards K.S., (2005). Flood frequency analysis: assumption and alternatives.
825 *Progress in Physical Geography*, 29(3), 392-410.
- 826 [38] Kite G.W., (1988). *Frequency and risk analysis in water resources*, Water Resources
827 Publications, Littleton, CO
- 828 [39] Kong X.M., Huang G.H., Fan Y.R., Li Y.P., (2015). Maximum entropy-Gumbel-Hougaard
829 copula method for simulation of monthly streamflow in Xiangxi river, China. *Stochastic
830 Environmental Research and Risk Assessment*, 29 (3), 915-927

- 831 [40] Lee T, Salas JD. 2011. Copula-based stochastic simulation of hydrological data applied to
832 Nile River flows. *Hydrology Research* 42(4): 318–330.
- 833 [41] Liu M., Xu X., Sun A.Y., Wang K., Liu W., Zhang X., (2014). Is southwestern China
834 experiencing more frequent precipitation extremes? *Environmental Research Letter*, 9,
835 064002.
- 836 [42] Ma M., Song S., Ren L., Jiang S., Song J., (2013). Multivariate drought characteristics using
837 trivariate Gaussian and Student copula. *Hydrological Processes*, 27, 1175-1190.
- 838 [43] Ma Z.Z., Wang Z.J., Xia T., Gippel C.J., Speed R., (2014). Hydrograph-based hydrologic
839 alteration assessment and its application to the Yellow River. *Journal of Environmental*
840 *Informatics*, 23(1), 1-13.
- 841 [44] Massey J.F., (1951) The Kolmogorov-Smirnov test for goodness of fit. *Journal of American*
842 *Statistical Association* 46 (253):68-78.
- 843 [45] Miao D.Y., Huang W.W., Li Y.P., Yang Z.F., (2014). Planning Water Resources Systems under
844 Uncertainty Using an Interval-Fuzzy De Novo Programming Method, *Journal of*
845 *Environmental Informatics*, 24(1), 11-23.
- 846 [46] Nelsen R.B., (2006). *An Introduction to Copulas*. Springer: New York.
- 847 [47] Razali N.M., Wah Y.B. (2011), Power comparisons of Shapiro-Wilk, Kolmogorov-Smirnov,
848 Lilliefors and Anderson-Darling tests. *Journal of Statistical Modeling and Analytics* 2(1): 21-
849 33
- 850 [48] Reddy J.M., Ganguli P., (2012). Bivariate flood frequency analysis of upper Godavari river
851 flows using Archimedean copulas. *Water Resources Management*, 26(14), 3995-4018.
- 852 [49] Rosenblatt M (1952) Remarks on a multivariate transformation. *The Annals of Mathematical*
853 *Statistics* 23(3):470-472.
- 854 [50] Reynolds D. (2007). Gaussian mixture models, *Encyclopedia of Biometric Recognition*
- 855 [51] Salvadori G, De Michele C (2004) Analytical calculation of storm volume statistics involving
856 Pareto-like intensity-duration marginals. *Geophysical Research Letters*, 31, L04502,
857 doi:10.1029/2003GL018767.
- 858 [52] Salvadori G. and C. De Michele. Multivariate real-time assessment of droughts via Copula-
859 based multi-site Hazard Trajectories and Fans. *Journal of Hydrology*, 2015. (in press).
- 860 [53] Salvadori, G., De Michele, C. 2004. Frequency analysis via copulas: Theoretical aspects and
861 applications to hydrological events. *Water Resources Research*, 40, 12, W12511.
- 862 [54] Salvadori, G., De Michele, C. 2010. Multivariate multiparameter extreme value models and
863 return periods: a copula approach. *Water Resources Research*, 46, 10, W10501.
- 864 [55] Salvadori G., De Michele C., Durante F., (2011). On the return period and design in a
865 multivariate framework. *Hydrology and Earth System Sciences* 15: 3293–3305
- 866 [56] Salvadori G., De Michele C., Kottegoda N.T., Rosso R., (2007). *Extremes in Nature: an*
867 *Approach using Copula*. Springer: Dordrecht; 292

- 868 [57] Salvadori, G., Durante F., De Michele, C. 2013. Multivariate return period calculation via
869 survival functions. *Water Resources Research*, 49, 2308–2311
- 870 [58] Samaniego, L., A. Bárdossy, and R. Kumar (2010), Streamflow prediction in ungauged
871 catchments using copula-based dissimilarity measures, *Water Resources Research*, 46,
872 W02506, doi:10.1029/2008WR007695.
- 873 [59] Serinaldi, F. (2009), A multisite daily rainfall generator driven by bivariate copula-based
874 mixed distributions, *Journal of Geophysical Research - Atmospheres*, 114,
875 doi:10.1029/2008JD011258.
- 876 [60] Shiau JT (2006) Fitting drought duration and severity with two-dimensional copulas. *Water*
877 *Resources Management*, 20, 795–815, 2006.
- 878 [61] Singh V. P., Sinclair R. A. Two-distribution method for flood frequency analysis. *Journal of*
879 *Hydraulics Division*, 1972, 98(1), 29-44.
- 880 [62] Singh V. P., Wang S. X., Zhang L. Frequency analysis of nonidentically distributed
881 hydrologic flood data. *Journal of Hydrology*, 2005, 307(1-4), 175-195.
- 882 [63] Sondergaard T. and Lermusiaux, P.F.J. (2013). Data assimilation with Gaussian mixture
883 models using the dynamically orthogonal field equations. Part I: Theory and scheme.
884 *Monthly Weather Review*, 141, 1761–1785
- 885 [64] Song S., Singh V.P., (2010). Meta-elliptical copulas for drought frequency analysis of
886 periodic hydrologic data. *Stochastic Environmental Research and Risk Assessment*, 24(3),
887 425-444.
- 888 [65] Sraj M., Bezak N., Brilly M. (2014). Bivariate flood frequency analysis using the copula
889 function: a case study of the Litija station on the Sava River. *Hydrological Processes*, DOI:
890 10.1002/hyp.10145
- 891 [66] Vandenberghe S, Verhoest NEC, De Baets B. (2010). Fitting bivariate copulas to the
892 dependence structure between storm characteristics: a detailed analysis based on 105 year 10
893 min rainfall. *Water Resources Research*, 46. DOI: 10.1029/2009wr007857.
- 894 [67] Wang L.Z., Huang Y.F., Wang L., Wang G.Q., (2014). Pollutant flushing characterization of
895 stormwater runoff and their correlation with land use in a rapidly urbanizing watershed.
896 *Journal of Environmental Informatics*, 23(1), 37-43.
- 897 [68] Waylen P., Woo M. K. Prediction of annual floods generated by mixed processes. *Water*
898 *Resources Research*, 1982, 18(4), 1283-1286.
- 899 [69] Willmott CJ, Matsuura K (2005) Advantages of the mean absolute error (MAE) over the root
900 mean square error (RMSE) in assessing average model performance. *Climate Research*
901 30(1):79.
- 902 [70] Wu Z.Y., Lu G.H., Liu Z.Y., Wang J.X., Xiao H., (2013). Trends of extreme flood events in
903 the Pearl river basin during 1951-2010. *Advances in Climate Change Research*, 4(2), 110-116
- 904 [71] Xu H., Taylor R.G., Kinston D.G., Jiang T., Thompson J.R., Todd M.C., (2010). Hydrological
905 modeling of River Xiangxi using SWAT2005: A comparison of model parameterizations using

906 station and gridded meteorological observations. *Quaternary International*, 226, 54-59.

907 [72] Yang W., Yang Z.F., (2014). Evaluation of Sustainable Environmental Flows Based on the
908 Valuation of Ecosystem Services: a Case Study for the Baiyangdian Wetland, China. *Journal*
909 *of Environmental Informatics*, 24(2), 90-100

910 [73] Yen B.C., (1970). Risk Analysis in design of engineering projects. *Journal of Hydrologic*
911 *Engineering*, 96(4), 959-966.

912 [74] Yin H.F., Li C.A., (2001). Human impact on floods and flood disasters on the Yangtze River.
913 *Geomorphology*, 41(2-3), 105-109

914 [75] Yue S., (2001). A bivariate gamma distribution for use in multivariate flood frequency
915 analysis. *Hydrological Processes*, 15(6), 1033-1045

916 [76] Yue S., T.B.M.J., Ouarda, B. Bobee, P. Legendre, P. Bruneau, (1999). The Gumbel mixed
917 model for flood frequency analysis. *Journal of Hydrology*, 226, 88-100.

918 [77] Zhang L., (2005). Multivariate hydrological frequency analysis and risk mapping. Ph.D
919 dissertation, Louisiana State University. Baton Rouge, Louisiana (USA).

920 [78] Zhang L., Singh V.P., (2006). Bivariate flood frequency analysis using the copula method.
921 *Journal of Hydrologic Engineering*, 11, 150-164.

922 [79] Zhang L, Singh VP (2007) Bivariate rainfall frequency distributions using Archimedean
923 copulas. *Journal of Hydrology* 332(1-2):93-109

924 [80] Zhang L, Singh VP (2012). Bivariate rainfall and runoff analysis using entropy and copula
925 theories. *Entropy* 14:1784-1812.

926 [81] Zhang N., Li Y.P., Huang W.W., Liu J., An Inexact Two-Stage Water Quality Management
927 Model for Supporting Sustainable Development in a Rural System. *Journal of Environmental*
928 *Informatics*, 24(1), 52-64.

929 [82] Zhang Y., Li H., Yang D., (2012a). Simultaneous estimation of relative permeability and
930 capillary pressure using ensemble-based history matching techniques. *Transport in porous*
931 *media* 94 (1), 259-276.

932 [83] Zhang Y, Song C, Zheng S, Yang D. (2012b). Simultaneous estimation of relative
933 permeability and capillary pressure for tight formations from displacement experiments. SPE
934 paper 162663, *the SPE Canadian Unconventional Resources Conference*; 2012 October 30-
935 November 1; Calgary, AB

936 [84] Zhang Y., Yang D., (2013). Simultaneous estimation of relative permeability and capillary
937 pressure for tight formations using ensemble-based history matching method. *Computers &*
938 *Fluids* 71, 446-460

939 [85] Zhang Y., Yang D., (2014). Estimation of relative permeability and capillary pressure for tight
940 formations by assimilating field production data. *Inverse Problems in Science and*
941 *Engineering* 22 (7), 1150-1175

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944 **Captions of Figures**

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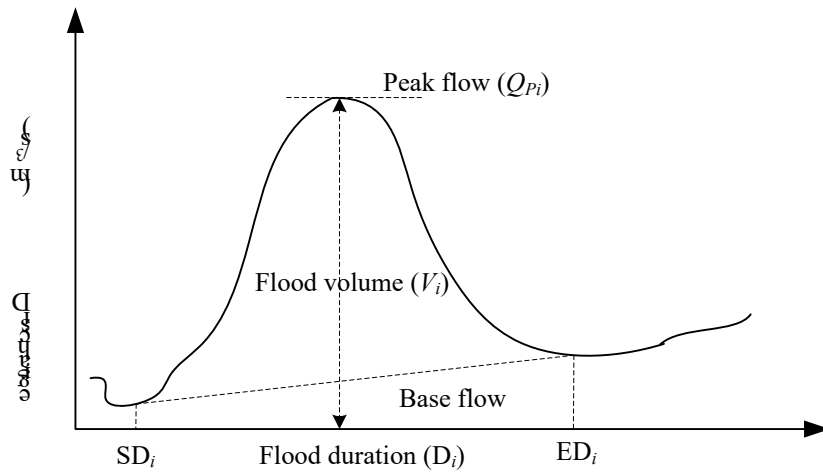
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Figure 1: the location of the studied watershed



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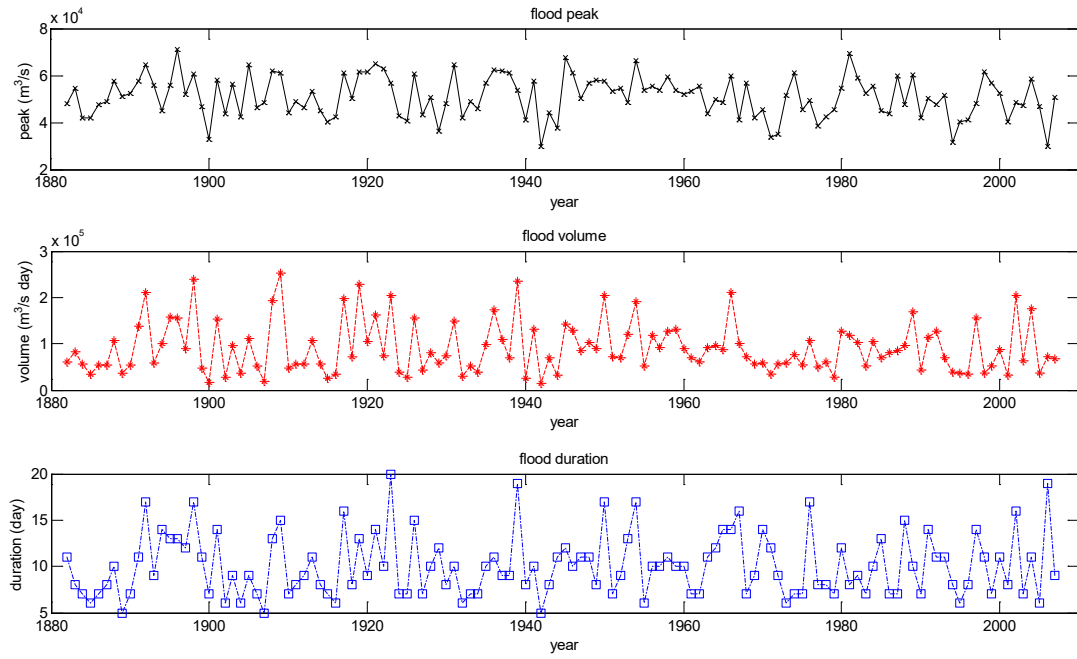
979 Figure 2. Typical flood hydrograph showing flood flow characteristics (adapted from

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[Ganguli and Reddy, \[22\]](#))

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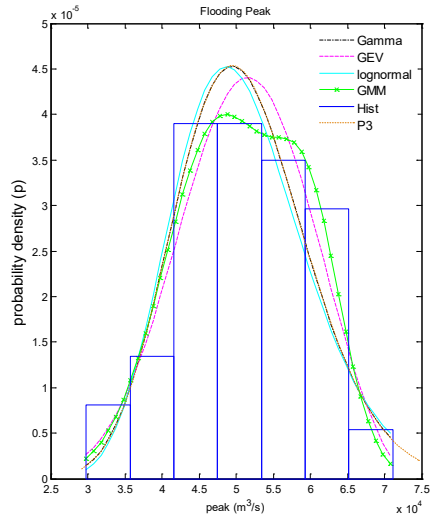
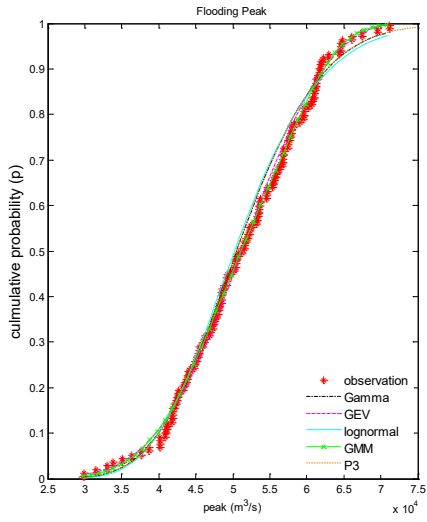


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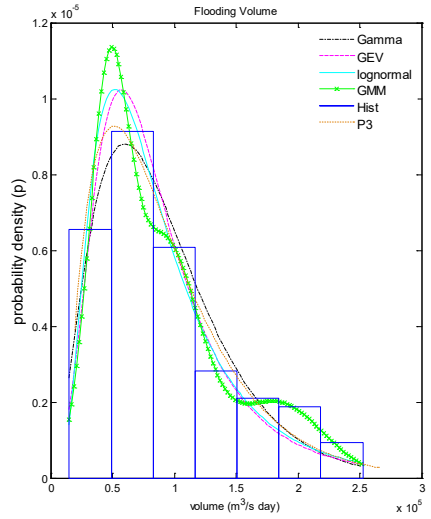
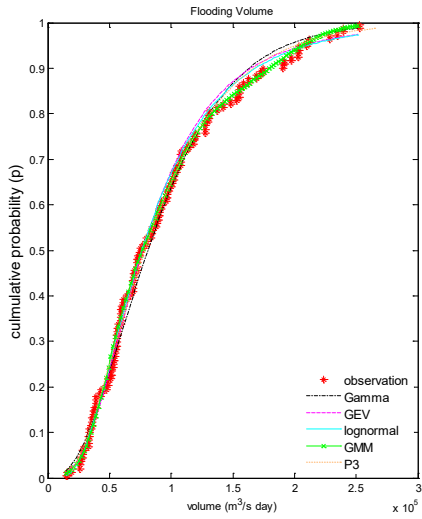
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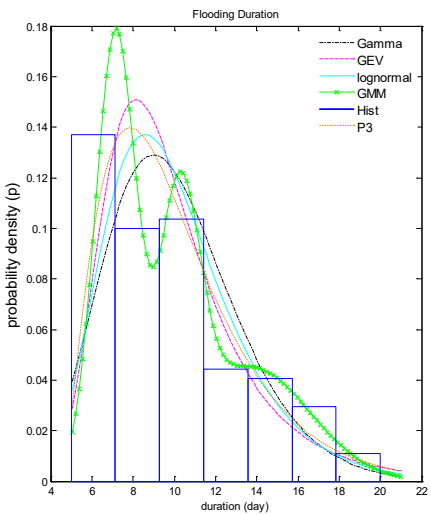
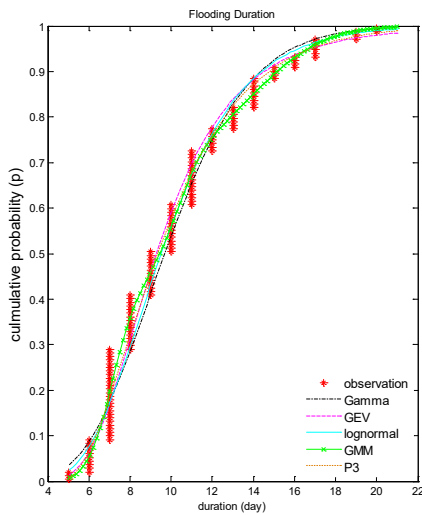
Figure 3. Variations of flood variables during the studied period



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Figure 4. Comparison of different probability density estimates with observed frequency.

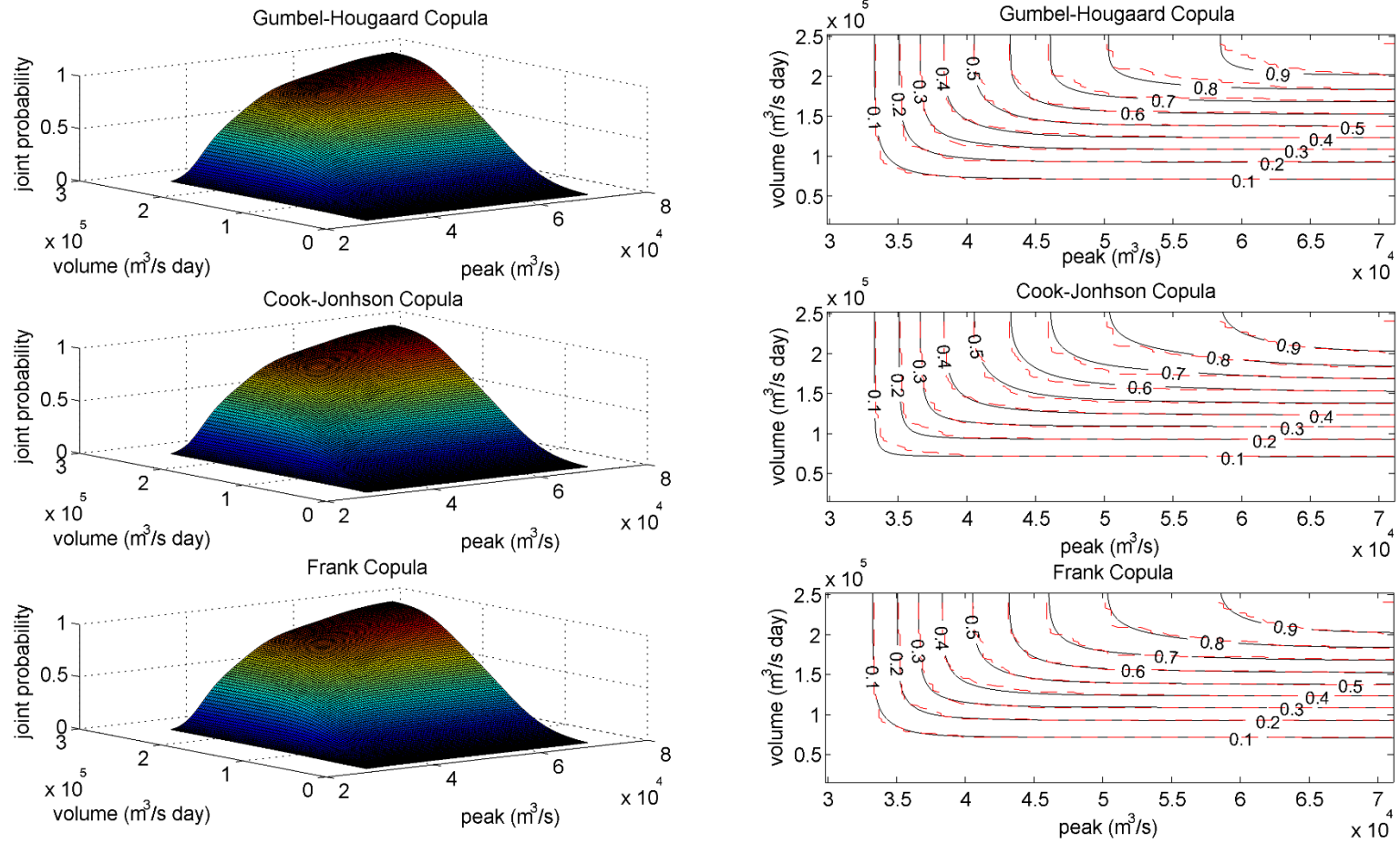


Figure 5. The copula estimation between flood peak and volume

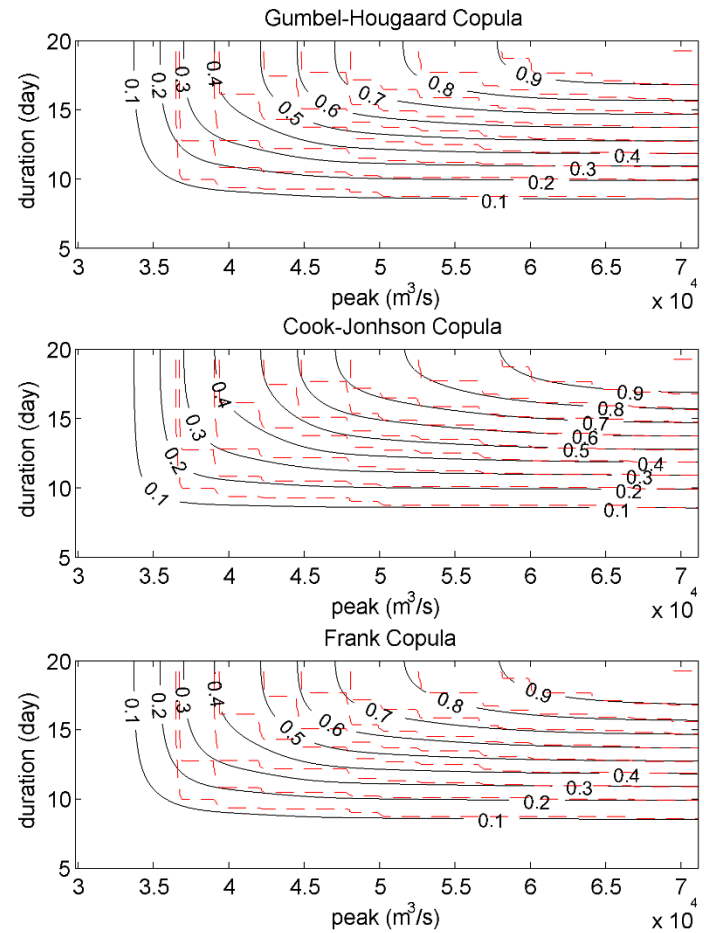
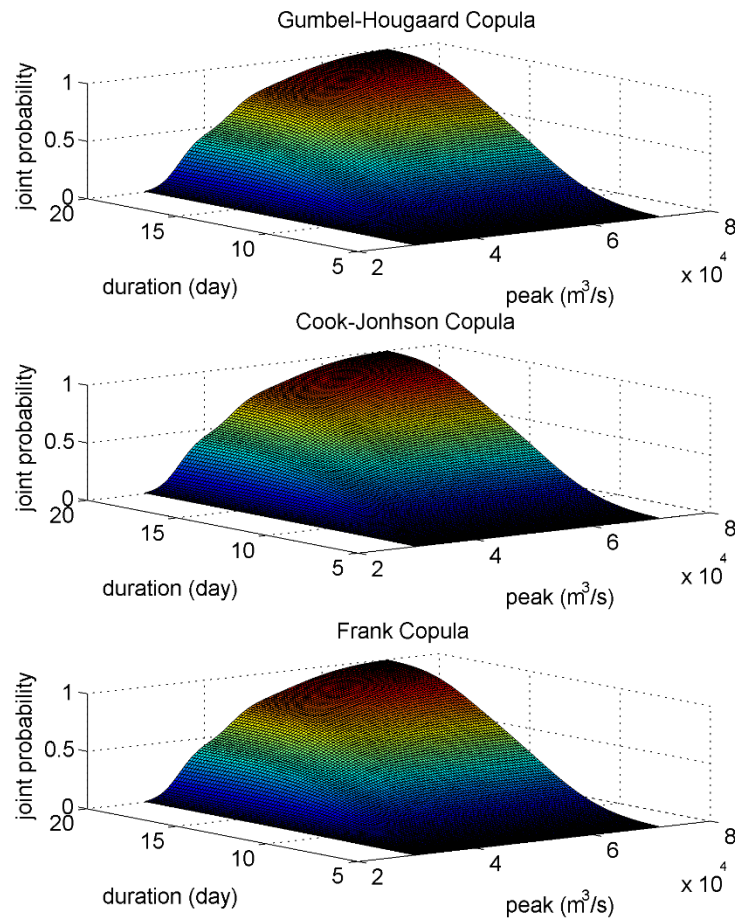


Figure 6. The copula estimation between flood peak and duration

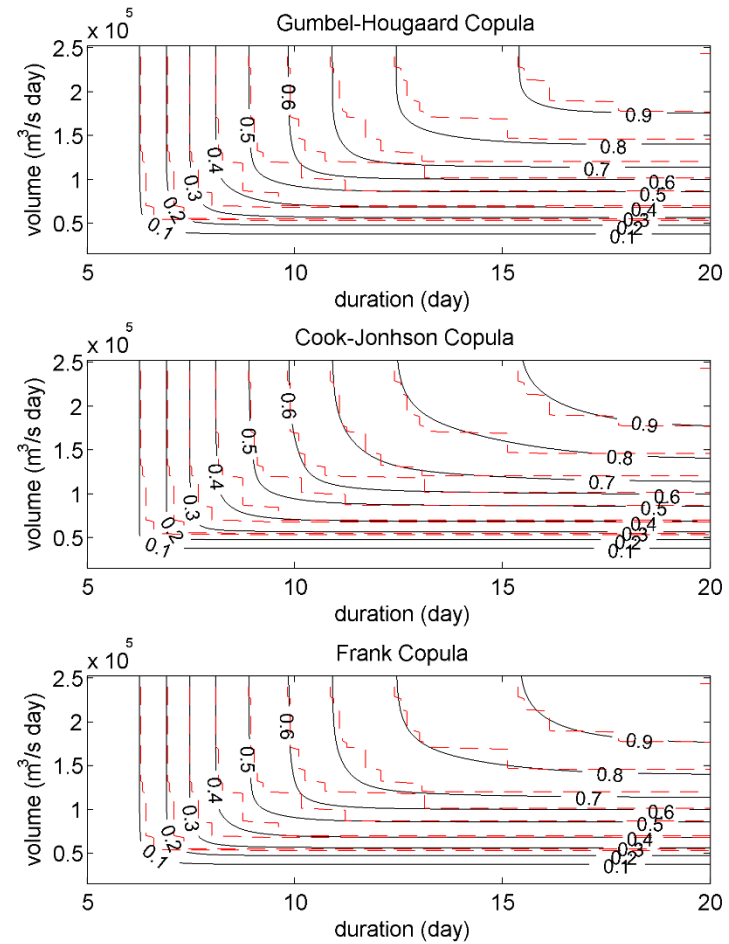
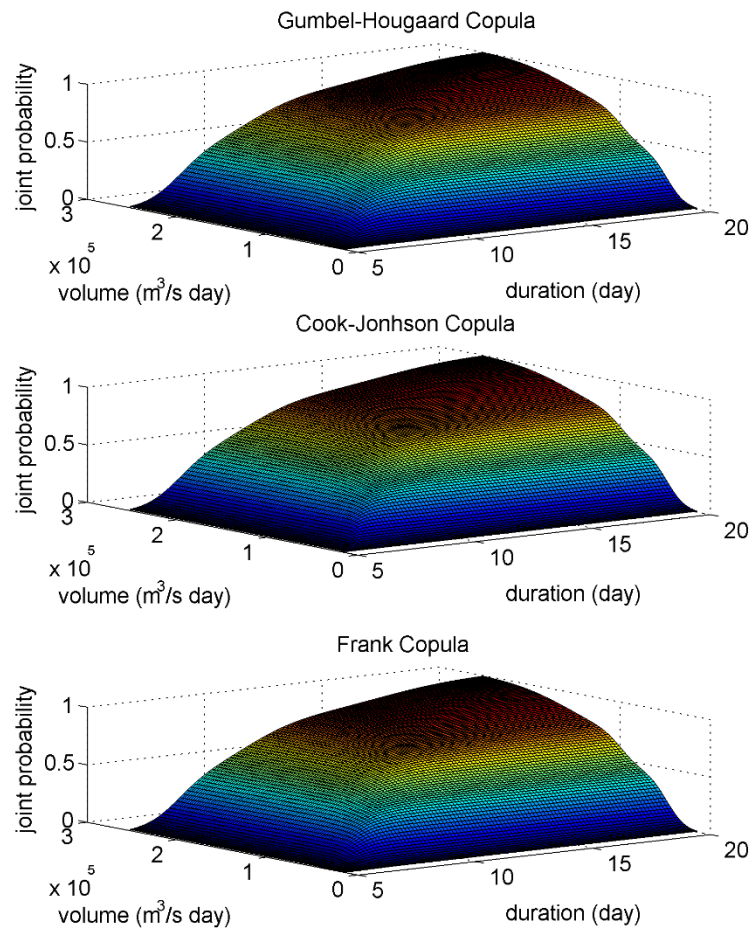


Figure 7. The copula estimation between flood volume and duration

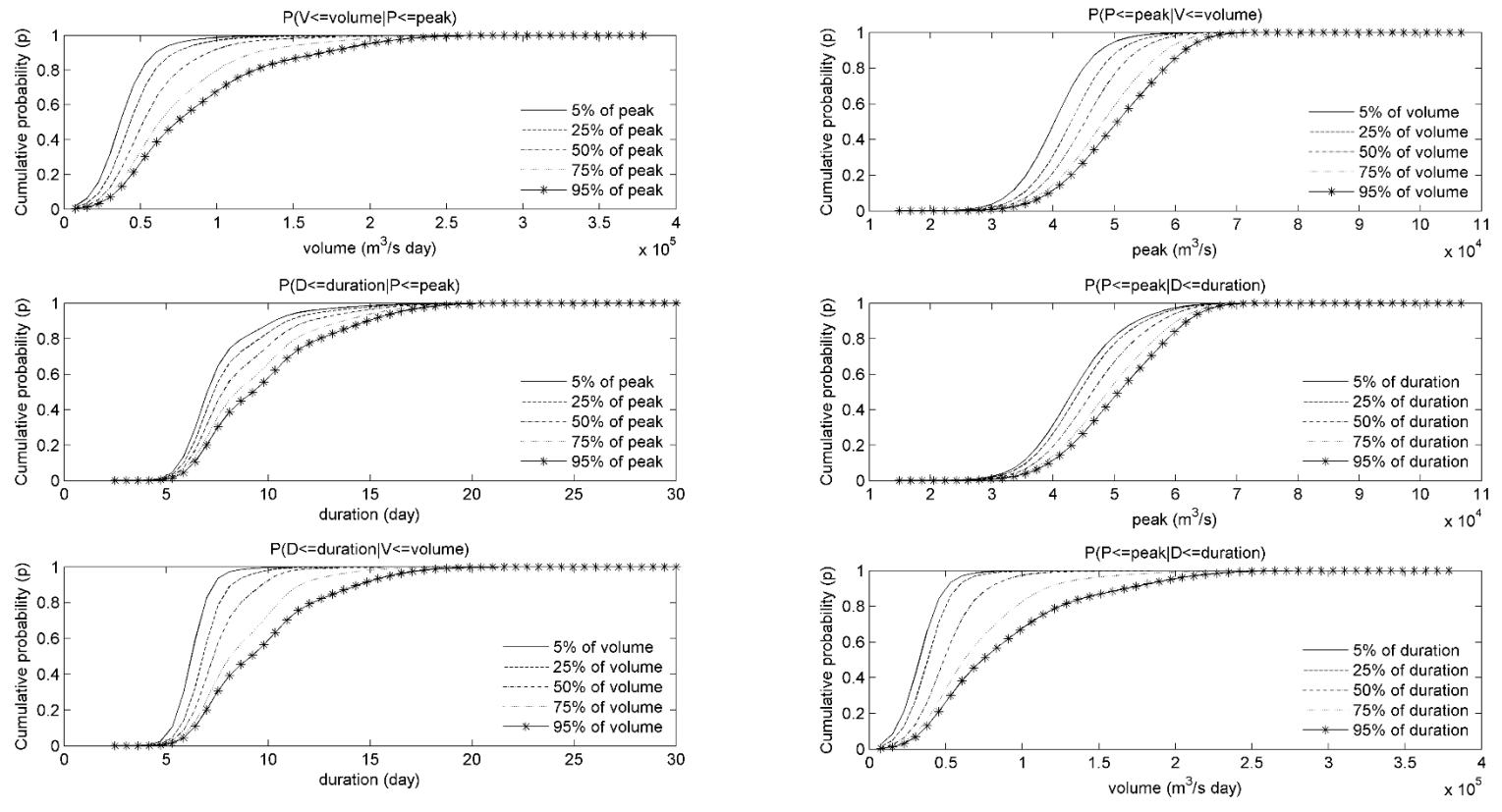


Figure 8. The conditional cumulative distribution functions.

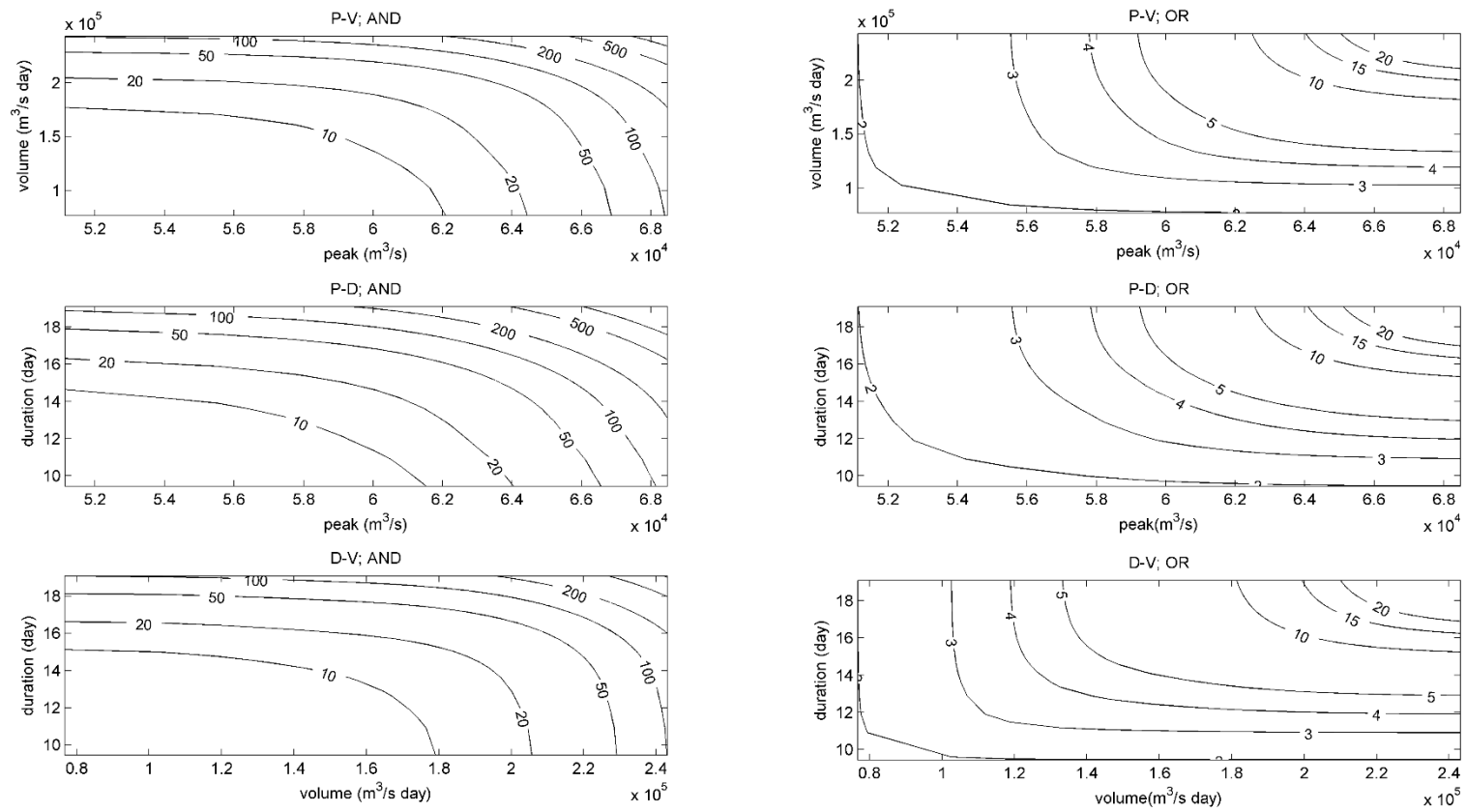


Figure 9. Comparison of the joint return periods.

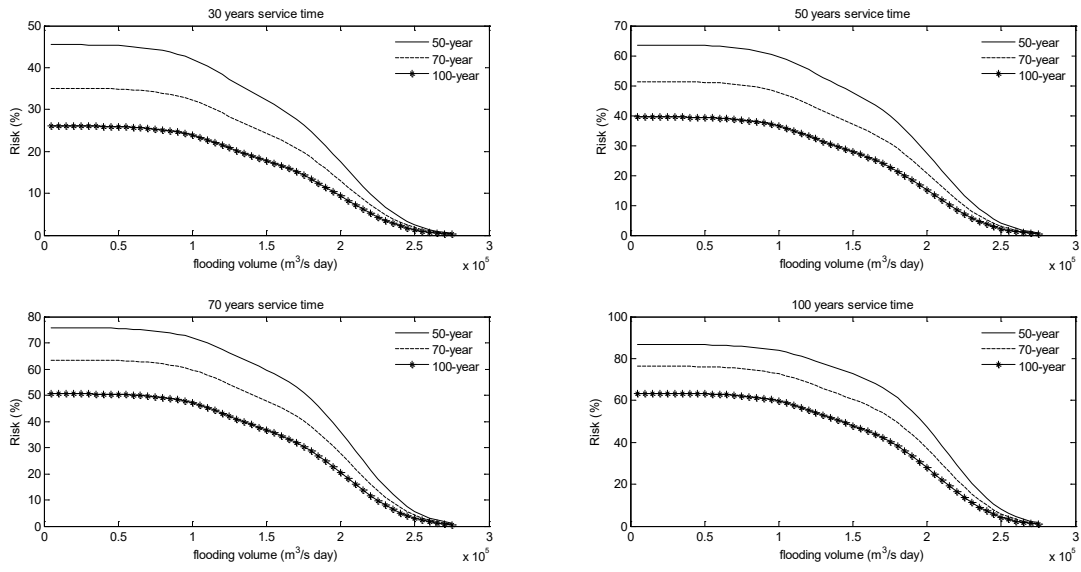


Figure 10. Bivariate flood risk under different flood peak-volume scenarios

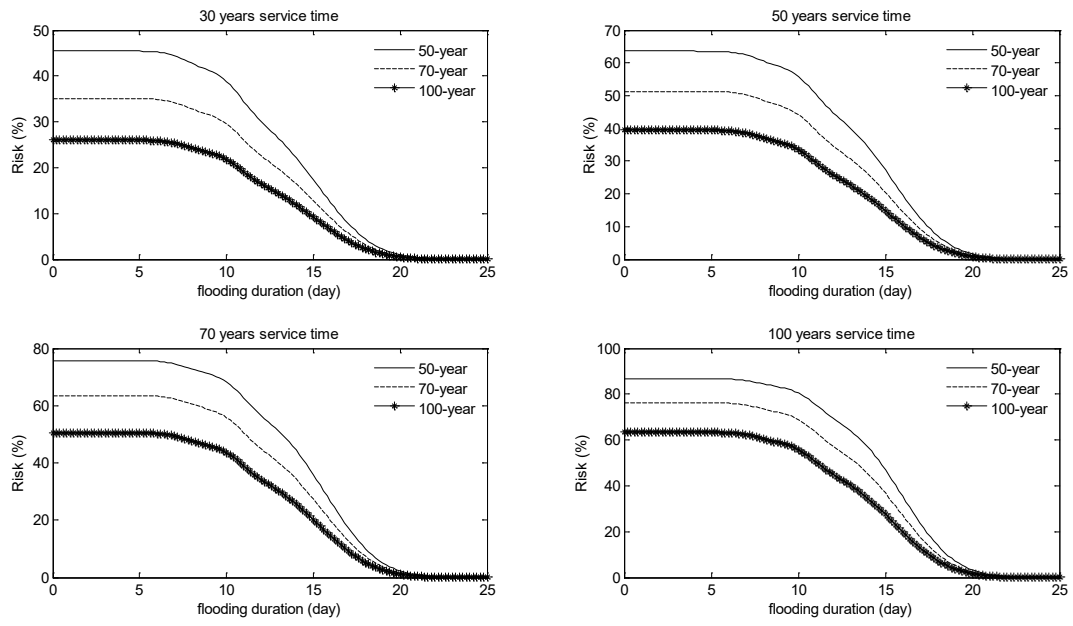


Figure 11. Bivariate flood risk under different flood peak-duration scenarios

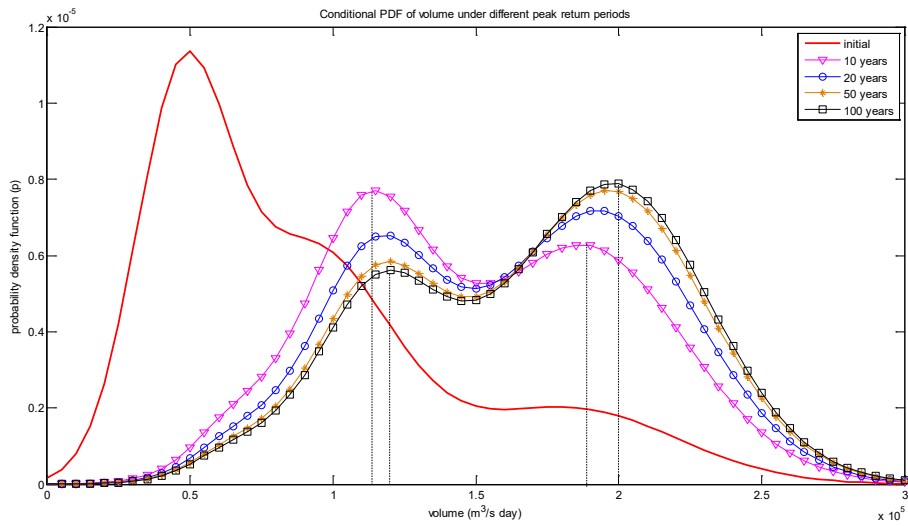


Figure 12. Probability density functions of volume under different peak flow return periods.

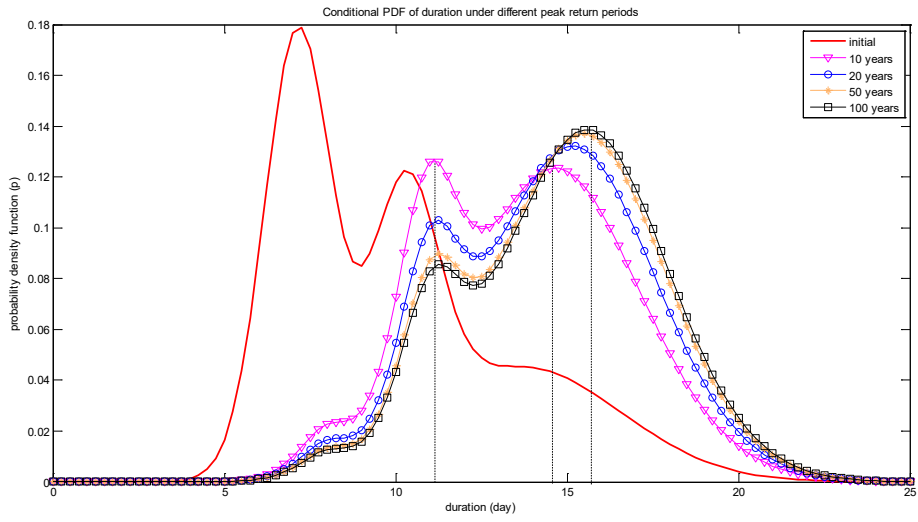


Figure 13. Probability density functions of duration under different peak flow return periods.

Table 1. Basic properties of applied copulas

Copula Name	Function[$C_\theta(u_1, u_2)$]	$\theta \in$	Generating function [$\phi(t)$]	$\tau = 1 + 4 \int_0^1 \frac{\phi(t)}{\phi'(t)} dt$
Cook- Johnson	$[u_1^{-\theta} + u_2^{-\theta} - 1]^{-1/\theta}$	$[-1, \infty) \setminus \{0\}$	$t^{-\theta} - 1$	$\frac{\theta}{\theta + 2}$
Gumbel-Hougaard	$\exp\{-[(-\ln u_1)^\theta + (-\ln u_2)^\theta]^{1/\theta}\}$	$[1, \infty)$	$(-\ln t)^\theta$	$1 - \theta^{-1}$
Frank	$-\frac{1}{\theta} \ln\left\{1 + \frac{(e^{-\theta u} - 1)(e^{-\theta v} - 1)}{e^{-\theta} - 1}\right\}$	$[-\infty, \infty) \setminus \{0\}$	$\ln\left[\frac{e^{-\theta t} - 1}{e^{-\theta} - 1}\right]$	$1 - \frac{4}{\theta} [D_1(-\theta) - 1]^*$

Note: * D_1 is the first order Debye function, and for any positive integer k , $D_k(x) = \frac{k}{x^k} \int_0^k \frac{t^k}{e^t - 1} dt$

Table 2. Parameters of marginal distribution functions of flood variables

Name	Probability density function	Parameters			
		Peak	Volume	Duration	
Gamma	$\frac{1}{b^a \Gamma(a)} x^{a-1} e^{-\frac{x}{b}}, \Gamma(a) = \int_0^\infty u^{a-1} e^{-u} du$	a	32.76	2.9	7.90
		b	1557.5	31363.5	1.24
GEV	$(\frac{1}{\sigma}) \exp(- (1+k \frac{(x-\mu)}{\sigma})^{\frac{1}{k}}) (1+k \frac{(x-\mu)}{\sigma})^{-1-\frac{1}{k}}$	k	-0.336	0.18	0.04
		μ	8899.6	36511.23	2.74
		σ	48177	63161.81	8.07
Lognormal	$\frac{1}{x \sigma_y \sqrt{2\pi}} \exp(-\frac{(y-\mu_y)^2}{2\sigma_y^2})$ $y = \log(x), x > 0, -\infty < \mu_y < \infty, \sigma_y > 0$	μ_y	10.82	11.24	2.22
		σ_y	0.18	0.62	0.36
Pearson Type III	$\frac{1}{b^a \Gamma(a)} (x-\alpha)^{a-1} e^{-\frac{x-\alpha}{b}}, \Gamma(a) = \int_0^\infty u^{a-1} e^{-u} du$	a	32.85	1.98	2.48
		b	1554.1	40002.4	2.22
		α	-21.91	12249.8	4.57

Table 3. Marginal distributions for flood variables through GMM

Flood Variables	Weights	Mean	Standard Deviation
Volume	0.4232	91987.0	27586.0
	0.1882	182387.1	37581.9
	0.3886	46691.3	16094.4
Peak	0.7436	47928	7551.2
	0.2564	60020	4480.5
Duration	0.2681	5.98	0.7
	0.4533	9.43	1.7
	0.2785	14.03	2.8

Table 4. Statistical test results for marginal distribution estimation

Flood variables	Marginal distribution	K-S test		RMSE	AIC
		<i>T</i>	P-value		
Peak	Gamma	0.0570	0.4253	0.0247	-401.0
	GEV	0.0362	0.7017	0.0176	-436.1
	Lognormal	0.0612	0.3740	0.0287	-384.6
	Pearson Type III	0.0610	0.7369	0.0246	-399.5
	GMM	0.0380	0.6776	0.0119	-473.0
Volume	Gamma	0.0611	0.3753	0.0266	-392.9
	GEV	0.0459	0.5705	0.0213	-415.2
	Lognormal	0.0390	0.6648	0.0174	-439.4
	Pearson Type III	0.0417	0.9808	0.0166	-442.5
	GMM	0.0434	0.6049	0.0148	-443.1
Duration	Gamma	0.1009	0.0716	0.0375	-355.3
	GEV	0.1023	0.0666	0.0403	-345.5
	Lognormal	0.0996	0.0769	0.0376	-355.1
	Pearson Type III	0.1113	0.0881	0.0378	-352.5
	GMM	0.0703	0.2754	0.0297	-366.9

Table 5. Dependence evaluations among flood variables

No.	Flood characteristics	Kendall's tau	Pearson's r
1	Peak – Volume	0.5509	0.6598
2	Volume – Duration	0.6756	0.7529
3	Peak - Duration	0.3561	0.2902

Table 6. Statistical test results for the flood pairs of peak-volume, peak-duration and volume-duration

Site	Copulas	Cramér von Mises statistic		RMSE
		S_n	P-value	
Peak - Volume	G-H	70.8224	0.3365	0.0168
	C-J	69.3597	0.3085	0.0199
	Frank	70.3817	0.3495	0.0149
Peak - Duration	G-H	66.2142	0.1215	0.0349
	C-J	64.9940	0.1165	0.0342
	Frank	65.7948	0.1325	0.0334
Volume-Duration	G-H	77.2530	0.1156	0.0302
	C-J	75.8958	0.1096	0.0305
	Frank	76.8450	0.1216	0.0291

Table 7. Comparison of univariate, bivariate return periods for flood characteristics (year)

T	Peak (m ³ /s)	Volume (m ³ /s day)	Duration (day)	T_{PV}^{AND}	T_{PD}^{AND}	T_{DV}^{AND}	T_{PV}^{OR}	T_{PD}^{OR}	T_{DV}^{OR}	\bar{T}_{PV}	\bar{T}_{PD}	\bar{T}_{DV}
5	59120.7	132815.5	12.9	8.5	11.4	7.2	3.5	3.2	3.8	6.4	7.7	5.9
10	62281.3	179597.7	15.2	24.5	36.6	19.2	6.3	5.8	6.8	16.2	22.0	13.8
20	64567.8	205920.7	16.6	78.8	127.6	58.0	11.5	10.8	12.1	46.8	70.8	36.7
50	66950.9	229240.1	18.1	419.6	726.4	290.3	26.6	25.9	27.4	227.3	380.1	163.1
100	68475.7	243091.0	19.1	1579.9	2809.4	1063.1	51.6	50.9	52.5	824.1	1438.0	566.2

Table 8. Statistical characteristics of the conditional PDFs of flood duration and volume under different peak flow return periods.

Flood variables	Index	initial	Return periods of peak flow (year)			
			10	20	50	100
Volume	Mean	91356.3	151750.3	161656.9	167571.8	169531.0
	Std	54840.2	51379.1	51441.9	51214.0	51095.1
	Kurtosis	0.4	-0.7	-0.8	-0.7	-0.7
	Skewness	1.0	0.1	-0.1	-0.2	-0.2
Duration	Mean	10.0	13.7	14.3	14.6	14.7
	Std	3.4	3.0	3.0	3.0	2.9
	Kurtosis	0.1	-0.4	-0.4	-0.3	-0.3
	Skewness	0.8	0.1	0.0	-0.1	-0.1